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Modeling

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Chromium spills on soils are causing severe groundwater contamination problems. To increase the performance of cleanup methods, the chemical behavior of chromium in soils and its transport through soils have to be understood. The objectives of this research were to investigate the sorption characteristics of chromium(VI) in soil and to develop a solute transport model that allows the prediction of chromium(VI) movement through soil columns. Of particular interest was the sorption behavior of chromium(VI) in the presence of phosphate, another adsorbing anion. Laboratory experiments were performed using soil samples taken from an area close to a chromium contamination site. Batch experiments were conducted for both chromium(VI) and phosphates to determine sorption parameters. Soil column breakthrough curves for chromium(VI) were determined with and without phosphate present. The batch studies indicated a strong kinetic sorption behavior for both chromium(VI) and phosphate, which was attributed to a physical

non-equilibrium transport process. The soil column experiments showed that the chromium(VI) removal rate from soil can be increased when phosphates are introduced into the soil solution during the desorption process. To interpret the results of these experiments, a physical non-equilibrium solute transport model was developed that incorporated competitive sorption. Numerical solutions of the transport equations were obtained by a partially implicit finite difference method. Computer simulations of experimental breakthrough and desorption curves had close agreement with experimental results. This study demonstrated that phosphate addition to the soil solution is a possible way of increasing desorption rates and, consequently, the effectiveness of chromium(VI) removal from contaminated soils.

# Chromium(VI) Sorption in Soils: Chemical Behavior and Solute Transport Modeling

by

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Dedicated to Traudl for her love and support

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### CHROMIUM(VI) SORPTION IN SOILS: CHEMICAL BEHAVIOR AND SOLUTE TRANSPORT MODELING

#### INTRODUCTION

Groundwater is an important freshwater resource. Over the last decade, more and more cases of groundwater contamination by heavy metals, in particular chromium, have been discovered. These contaminations are of great concern due to the high toxicity exhibited by Cr(VI), posing a severe threat to human health as well as to the environment. Efficient cleanup methods are therefore required to preserve aquifers as a valuable freshwater source. To increase the cleanup efficiency it is necessary to be able to predict the movement of chromium through soils and aquifer systems and thus understand its transport mechanisms, its chemical behavior and its interactions with the soil matrix.

The objectives of this work were to investigate chromium(VI) behavior in a natural soil system, in particular to determine the effect of pH on Cr(VI) sorption onto the soil and to better understand the kinetic effects that accompany chromium(VI) sorption. Another important objective was to investigate the Cr(VI) removal effectiveness from contaminated soil when phosphate is used as a competitive anion during chromium(VI) desorption. As a step to improve the understanding of the chromium(VI) movement in aquifer systems, a solute transport model had to be developed to describe the movement of chromium(VI) through laboratory soil columns. This model had to be verified using experimentally determined breakthrough curves (BTC's).

#### LITERATURE REVIEW

Chromium(VI) adsorption studies have been conducted using a variety of solid media. Griffin et al.(9) conducted an adsorption study on kaolinite and montmorillonite clays. They found that Cr(VI) adsorption decreased as pH increased. Below pH 2, however, Cr(VI) adsorption also decreased gradually. These results were partially attributed to an increase of positive surface charges on clavs and hydrous oxides as the pH was lowered. It was also believed that Cr(VI) speciation played an important role. At low pH, the fraction of HCrO<sub>4</sub> present becomes smaller, favoring the uncharged H<sub>2</sub>CrO<sub>4</sub>. This was believed to be the reason for less Cr(VI) adsorption below pH 2. At high pH, when Cr(VI) is present as CrO<sub>4</sub><sup>2</sup>, no sorption was observed. It was believed that the double charge on the CrO<sub>4</sub><sup>2</sup> anion caused it to be repelled by the negative clay surface charges. Davis and Leckie (5) conducted a Cr(VI) adsorption study on amorphous iron oxyhydroxide. In this study it was concluded that both  $CrO_4^2$  and  $HCrO_4$  are sorbing species. The adsorption mechanism of Cr(VI) was also discussed by Stollenwerk and Grove (11). Here it was concluded that Cr(VI) adsorbed by nonspecific processes as well as by specific sorption site processes. The fraction of Cr(VI) that could be extracted easily from the solid phase was assumed to be adsorbed by nonspecific processes whereas the remainder of the Cr(VI) on the solid phase was assumed to be adsorbed by specific sorption site processes. The effect of other anions in solution on Cr(VI) adsorption was also investigated. It was found that the amount of chromium(VI) adsorbed is a function of the type

and concentration of the other anions in solution. A strong competitive effect observed for phosphate was explained by direct competition for specific surface sites. The smaller effect of Cl<sup>-</sup> and NO<sub>3</sub><sup>-</sup> on Cr(VI) adsorption was attributed to a decrease in electrostatic potential near the surface of a particle, leading to less anion adsorption by nonspecific processes. An earlier study by Bartlett and Kimble (2) showed similar results for phosphate, attributing its good extraction effect to specific site competition. A study of chromium(VI) adsorption on iron oxyhydroxides in the presence of paired solute systems and multiple ion mixtures was recently conducted by Zachara et al.(13). Their findings agreed with the previously conducted research, showing decreased Cr(VI) adsorption for each anion added to the solute mixture.

Another important mechanism for chromium(VI) behavior in natural soil systems, besides adsorption, is the transformation to Cr(III) by reduction. Bartlett and Kimble (2) found fast Cr(VI) reduction in soil in the presence of soil organic matter. An almost completely organic-free soil showed little Cr(VI) reduction potential. However, when manure was added and pH was adjusted below 3, most of the added Cr(VI) was reduced within 24 hours. In this study, no oxidation of Cr(III) to Cr(VI) was observed. In a later study, however, Bartlett and James (3) found Cr(III) oxidation in soil in the presence of oxidized manganese. A reduction process of Cr(VI) to Cr(III) followed by a precipitation of Cr(OH)<sub>3</sub> was assumed to be at least partially the reason for a poor Cr(VI) recovery efficiency observed by Stollenwerk and Grove (11) for their soil column experiments. Extensive reduction of Cr(VI) in topsoil was also observed by Bloomfield and

Pruden (4). Experimental results also indicated a strong pH dependence for the reduction process leading to increased reduction rates with decreased pH.

A variety of solute transport models for both organic and inorganic species have been proposed. Most of these models incorporate the same dispersion-advection transport processes. They differ, however, in the way that sorption / desorption reactions are modeled. A review of sorption models for reactive solutes in soil was given by Travis and Etnier (24). Both equilibrium and first-order kinetic models were discussed. Equilibrium models were originally developed to describe sorption isotherms for gases and organic solutes. Bar-Yosef (14) derived a sorption model based on a competitive Langmuir isotherm for ionic species and applied it successfully to pH-dependent zinc adsorption by soils. This model did not include the effect of the electric potential (Y) and the electrolyte concentration on adsorption. Bar-Yosef believed that the effect of  $\Psi$  was to some extent accounted for by the binding constants of the different ions. A transport model based on ion exchange was presented by Valocchi et al. (25). Sorption processes for ionic species were modeled employing the well known principle of ion exchange selectivity. Although successful for the cases considered, the authors realized that the ion selectivity coefficients are typically not constant but instead are a function of the sorbed phase concentration. These coefficients are also highly variable in natural soil systems, thus hard to estimate. Cederberg et al. (15) presented a solute transport model that incorporated ion exchange and surface complexation coupled with a chemical equilibrium model. Good agreement was

found between predicted and measured concentrations of cadmium, chloride and bromide in laboratory soil columns.

For many transport problems, the assumption of equilibrium sorption is not valid (Van Genuchten (29)). Two classes of non-equilibrium models have been developed: chemical non-equilibrium and physical non-equilibrium.

Derivations of these models are in Van Genuchten and Cleary (28) and Van Genuchten (29).

The chemical non-equilibrium model assumes that there are two types of sorption sites. For one type of site the sorption reaction is assumed to be a slow, kinetically controlled process. For the other type of site sorption is assumed to be a kinetically faster or even an instantaneous equilibrium process. Such a two-site model was recently used by Selim and Amacher (23) to model chromium(VI) transport through three different soils. Kinetic Langmuir equations were used to describe the sorption reactions. This approach was partially successful in describing data from miscible displacement experiments. Grove and Stollenwerk (17) modeled chromium(VI) movement through alluvial materials coated with iron-oxide and hydroxide. They used a one-site chemical non-equilibrium model based on a kinetic Langmuir equation. This model agreed better with experimental data than when sorption was assumed to be at equilibrium.

Physical non-equilibrium models usually divide the total water content of a soil into a mobile and an immobile region and assume that the kinetic sorption behavior is caused by a physical mechanism, typically a diffusional process across the immobile water layer to the solid surface. Wu and Gschwend (30) recently

presented evidence that the sorption kinetics of hydrophobic organic compounds on natural sediments are controlled by such an intraparticle diffusion.

These two major groups of non-equilibrium solute transport models were compared by Nkedi-Kizza et al. (20). When a linear sorption reaction was used, the analytical solutions for the two models are equivalent. These authors concluded that the difference between the two models is of little practical importance and suggest that the occurrence of specific processes can only be verified with microscopic measurements.

#### CHAPTER I:

# HEXAVALENT CHROMIUM CHEMISTRY IN SOILS: LABORATORY EXPERIMENTS

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#### ABSTRACT

Knowledge of the fate of chromium(VI) in natural soil and aquifer systems is of great importance due to the high toxicity exhibited by Cr(VI) to both humans and the environment. The objectives of this work were to investigate Cr(VI) behavior in a Dayton series clay, in particular pH and kinetic effects on adsorption reactions. Another objective was to investigate the degree to which phosphate is able to increase the Cr(VI) extraction performance from contaminated soil.

Batch reactor experiments were conducted for chromium(VI) and phosphate. The results indicated that reduction of Cr(VI) to Cr(III) and subsequent removal, besides Cr(VI) adsorption, was a major process in soil suspensions. The total Cr(VI) removal increased with decreasing pH and increasing initial Cr(VI)

concentration. The adsorption process itself consisted of two steps, an initial fast uptake followed by a slower, kinetically controlled uptake. This behavior was attributed to a physical non-equilibrium process.

Soil column experiments were conducted for Cr(VI) using both distilled water and a phosphate solution as extraction agents. The phosphate solution improved the Cr(VI) removal effectiveness by increasing the chromium(VI) desorption rate significantly. This effect was attributed to specific anion competition. The amount of Cr(VI) recovered, however, did not increase significantly when phosphate was used compared to using distilled water for Cr(VI) desorption.

Additional Word Index: Chromium(VI) chemistry, adsorption, physical non-equilibrium, phosphate competition.

#### INTRODUCTION

Hexavalent chromium contaminations in aquifer systems are of great concern due to the high environmental toxicity exhibited by Cr(VI). Efficient cleanup methods are required to preserve aquifers as a valuable freshwater resource. To increase the cleanup efficiency it is necessary to better understand the behavior of chromium(VI) in natural soil systems. The objectives of this work were to investigate chromium(VI) behavior in a natural soil system, particularly to determine the effect of pH on Cr(VI) sorption onto the soil and to better understand the kinetic effects that accompany chromium(VI) sorption. Another important objective was to investigate the Cr(VI) removal effectiveness from contaminated soil when phosphate is used as a competitive anion during chromium(VI) desorption.

Chromium(VI) adsorption studies have been conducted using a variety of solid media. Griffin et al.(9) conducted an adsorption study on kaolinite and montmorillonite clays. They found that Cr(VI) adsorption decreased as pH increased. Below pH 2, however, Cr(VI) adsorption also decreased gradually. These results were partially attributed to an increase of positive surface charges on clays and hydrous oxides as the pH was lowered. It was also believed that Cr(VI) speciation played an important role. At low pH, the fraction of HCrO<sub>4</sub> present becomes smaller, favoring the uncharged H<sub>2</sub>CrO<sub>4</sub>. This was believed to be the reason for less Cr(VI) adsorption below pH 2. At high pH, when Cr(VI) is present as CrO<sub>4</sub><sup>2-</sup>, no adsorption was observed. It was believed that the double

charge on the CrO<sub>4</sub><sup>2</sup> anion caused it to be repelled by the negative clay surface charges. Davis and Leckie (5) conducted a Cr(VI) adsorption study on amorphous iron oxyhydroxide. In this study it was concluded that both  $CrO_4^2$  and  $HCrO_4^2$  are sorbing species. The adsorption mechanism of Cr(VI) was also discussed by Stollenwerk and Grove (11). Here it was concluded that Cr(VI) adsorbed by nonspecific processes as well as by specific sorption site processes. The fraction of Cr(VI) that could be extracted easily from the solid phase was assumed to be adsorbed by nonspecific processes whereas the remainder of the Cr(VI) on the solid phase was assumed to be adsorbed by specific sorption site processes. The effect of other anions in solution on Cr(VI) adsorption was also investigated. It was found that the amount of chromium(VI) adsorbed is a function of the type and concentration of the other anions in solution. A strong competitive effect observed for phosphate was explained by direct competition for specific surface sites. The smaller effect of Cl and NO<sub>3</sub> on Cr(VI) adsorption was attributed to a decrease in electrostatic potential near the surface of a particle, leading to less anion adsorption by nonspecific processes. An earlier study by Bartlett and Kimble (2) showed similar results for phosphate, attributing its good extraction effect to specific site competition. A study of chromium(VI) adsorption on iron oxyhydroxides in the presence of paired solute systems and multiple ion mixtures was recently conducted by Zachara et al.(13). Their findings agreed with the previously conducted research, showing decreased Cr(VI) adsorption for each anion added to the solute mixture.

Another important mechanism for chromium(VI) behavior in natural soil

systems, besides adsorption, is the transformation to Cr(III) by reduction. Bartlett and Kimble (2) found fast Cr(VI) reduction in soil in the presence of soil organic matter. An almost completely organic-free soil showed little Cr(VI) reduction potential. However, when manure was added and pH was adjusted below 3, most of the added Cr(VI) was reduced within 24 hours. In this study, no oxidation of Cr(III) to Cr(VI) was observed. In a later study, however, Bartlett and James (3) found Cr(III) oxidation in soil in the presence of oxidized manganese. A reduction process of Cr(VI) to Cr(III) followed by a precipitation of Cr(OH)<sub>3</sub> was assumed to be at least partially the reason for a poor Cr(VI) recovery efficiency observed by Stollenwerk and Grove (11) for their soil column experiments. Extensive reduction of Cr(VI) in topsoil was also observed by Bloomfield and Pruden (4). Experimental results also indicated a strong pH dependence for the reduction process leading to increased reduction rates with decreased pH.

#### MATERIALS AND METHODS

#### Materials

The soil used in the experiments was a Dayton series silty clay collected from an uncontaminated field at the Corvallis Airport (7-15-87), near the United Chrome Products Superfund site (Ecology and Environment (6)). Previous investigations at that site showed that most of the chromium contamination is found in a zone of clayey silt extending from the surface to roughly six meters of depth; there was little variation between the soil in this zone and other surface soils in the area (Ecology and Environment (6)). The soils used in the present study should therefore be representative of much of the contaminated soil. The soil was taken from a zone of roughly 5 cm to 35 cm depth. Field conditions were such that no drying was required before further preparation. Large clumps were broken up and pebbles and large roots were removed by mechanical grinding in a "Dynacrush" soil grinder followed by passing the soil through a 2 mm mesh sieve. The ground, sieved soil was hand mixed and stored at field moisture and 4 °C until use.

All chemicals used in this study were ACS reagent grade. Glass and plasticware were soaked in 10% nitric acid and rinsed well with glass distilled water prior to all uses.

#### Experimental Procedures

The pH of the point-of-zero-charge (PZC) of the Dayton soil was determined by soil titrations performed in background electrolyte solutions. Soil-electrolyte suspensions were prepared with one gram of soil and 25 milliliters of solution. Solution ionic strengths were adjusted to 0.01 M and 0.10 M for two of the titrations by addition of KCl to glass distilled water; pH was adjusted by strong acid or base (HCl or KOH). In a third titration, no salt was added; ionic strength depended on acid or base strength only.

The organic matter content of the soil was determined by weight loss after ashing soil samples for 2 hours at 550 °C.

Batch reactor experiments were performed for both chromium(VI) and phosphate. Chromium(VI) and phosphate solutions were reacted with soil in 50 milliliter, screw-top plastic centrifuge tubes. Unless otherwise noted, 25 milliliters of solution were reacted with one gram of soil. All Cr(VI) and phosphate solutions were prepared with glass distilled water. Soil-suspensions were kept well mixed by continuous shaking in a 25 °C constant-temperature shaker bath. The headspace of the centrifuge tubes was air filled; no effort was made to control the O<sub>2</sub>(g) or CO<sub>2</sub>(g) content. Hexavalent chromium was added in solution as potassium dichromate; stock solutions were prepared at concentrations 1 gram per liter (0.192 M). Phosphate solutions were prepared from KH<sub>2</sub>PO<sub>4</sub>. Strong acid or base (HCl or KOH) was added to adjust pH between 2.5 and 9. After shaking for

the desired reaction time, the tubes were centrifuged at 10,000 rpm for 10 minutes and the supernatant solutions were passed through 0.45 µm Millipore filters. Samples were stored in plastic containers until analysis for pH, total Cr, and/or Cr(VI), or phosphate, respectively. Adsorbed chromium and phosphate were determined by difference from initial and final solution concentration.

Soil column experiments were conducted to more closely simulate field conditions of continuous hydraulic flow through a stationary porous medium. The soil columns utilized the same soil (uncontaminated) as those of the batch experiments described above, packed to approximate field density (about 1.1 g/cm³). Columns employed were borosilicate glass with a fritted porous support plate. The 25 cm long by 3.2 cm inside diameter columns were typically packed with 50 g of soil to a porosity of 35-40 %, yielding a soil column of about 5.7 cm height. Hydraulic flow through the column was also set to approximate field conditions (about 11 ml/h).

For these column experiments chromium(VI) was first adsorbed on the soil columns in a distilled water solution until equilibrium (column effluent equals column influent chromium concentration) was approximately achieved between the soil and applied chromium solution. Then chromium(VI) was extracted from the soil columns by application of various extraction media. Initial Cr(VI) concentrations, pH, and different extractants that were used in the various column studies are presented in Table I.1.

Another soil column experiment was performed using a chloride (Cl<sup>-</sup>)

solution. The soil column for this tracer experiment was prepared in the same manner as described above. The specific conditions of this experiment are also presented in Table I.1.

Table I.1 Soil column continuous flow experiments.

no	Pore volume ml	Initial Cr(VI) conc. mg/l	pН	Extractant
1	18.5	10	4.10	Distilled water
2	18.5	50	3.85	Distilled water
3	18.4	10	4.53	$H_2PO_4^- 0.02 M$
Cl tracer colu	ımn experimen	<b>t:</b>		
Evneriment	Pore volume	Initial Cl <sup>-</sup>	pН	Extractant
no	ml	conc. mg/l		

#### Analytical Methods

Chromium(VI) concentrations in solution were determined by the colorimetric method of Bartlett and Kimble (1). One milliliter of an s-diphenyl carbazide reagent solution was mixed with 1 ml of sample or standard plus 7 mls H<sub>2</sub>O and measured for absorbance at 540 nm on a Bausch and Lomb Spectronic 88 spectrophotometer. The reagent solution was prepared by dissolving 200 mg sdiphenyl carbazide in 100 ml of 95% ethanol and adding 120 mls 85% H<sub>3</sub>PO<sub>4</sub> in 280 mls distilled water. A small amount of KMnO<sub>4</sub> was added until a pink color developed: then the mixture was heated at 60 °C until the color disappeared. This solution was stored at 4 °C in a dark glass bottle. Standard solutions of 1, 3 and 5 mg/l were prepared from a 1000 mg/l K<sub>2</sub>Cr<sub>2</sub>O<sub>7</sub> stock solution. Concentrations of samples were calculated by using linear regression on the standard concentrations. The detection limit was determined to be 0.01 mg/l. Total chromium in solution was determined by flame atomic absorption spectroscopy using a Perkin Elmer Model 360 AAS. The detection limit was determined to be 1 µM (0.050 mg/l). The precision of this method was lower than for the colorimetric method; variations in results of up to 20% were sometimes observed in repeated analyses of individual samples in the optimum concentration range. Trivalent chromium (Cr(III)) concentrations were calculated by difference, subtracting the Cr(VI) from total chromium concentrations.

Phosphate and chloride concentrations were determined by anion chromatography using a DIONEX Series 4000i ion chromatograph (IC) equipped

with a conductivity detector. The columns used in the IC were a HPIC-AG4A guard column and a HPIC-AS4A separator column. The eluent solution injected was a 1.8 mM Na<sub>2</sub>CO<sub>3</sub> + 1.7 mM NaHCO<sub>3</sub> at a flow rate of 2.0 ml/min.

3 ml/min of 25 mN H<sub>2</sub>SO<sub>4</sub> solution were injected into the column as regenerant. Phosphate standards of 10, 20, 30 and 50 mg/l (as H<sub>2</sub>PO<sub>4</sub>) and chloride standards of 1, 5, 10 and 15 mg/l (as NaCl) were used to generate standard curves. Concentrations of unknowns were calculated by linear regression on the standard concentrations.

Solution pH was measured with an Orion research grade Ag/AgCl glass combination electrode (Model 91-02) with an Orion model 601a digital analyzer. Readings were made after five minute equilibration between electrode and sample solution. Before use, the system was calibrated to the appropriate pH range using buffer standard solutions prepared from METREPAK pHydiron buffer capsules.

#### RESULTS AND DISCUSSION

#### Soil Analyses

The pH of the point-of-zero-charge (PZC) of a soil is defined as the pH value at which the total net charge on the surface of a solid particle is zero (Sposito (10)). This value is commonly determined by soil titration experiments performed in background electrolyte solutions at two or more ionic strengths (Sposito (10); Stumm and Morgan (12)). When the net surface charge is zero, the activity of sorbed species is unaffected by ionic strength. Therefore, plots of proton or hydroxide surface densities versus pH intersect at the PZC. Figure I.1 shows the results for the soil used in this study. For each titration shown, some ions were dissolved from the soil into solution. However, no effort was made to quantify their contribution to the total ionic strength.

The determined PZC at pH 3.8 indicates that for a large pH range (pH>3.8) the soil contains more negative charges than positive charges and should electrostatically favor cation adsorption over anion adsorption. This could influence adsorption for anionic Cr(VI) as well as for cationic Cr(III). With a determined soil organic matter content of 4.4% and a relatively low soil pH, some Cr(VI) reduction to Cr(III) is likely to occur (Bartlett and Kimble (2)).

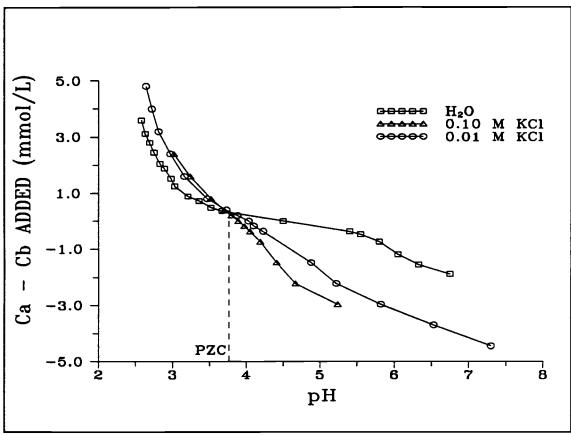


Figure I.1 Soil Point-of-Zero-Charge (PZC) determination using acid(Ca) - base(Cb) titrations in electrolytes with varying ionic strengths.

#### Cr(VI) Batch Reactor Experiments

A chromium(VI) adsorption batch reactor experiment was conducted to investigate the time dependence of the adsorption process for different pH values. Over the whole pH range considered, the adsorption reaction can be separated into two parts (Figure I.2). An initial, almost instantaneous uptake of chromium(VI) is followed by a much slower, kinetically controlled Cr(VI) uptake. The overall tendency is that the Cr(VI) uptake increases with decreasing pH. This agrees with the findings of Griffin et al.(9) and was explained by an increase in positive

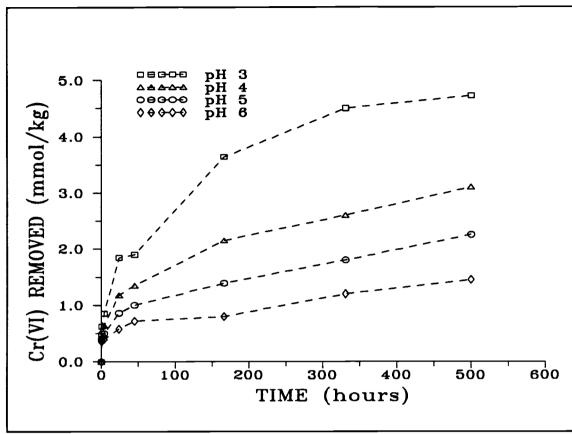


Figure I.2 Chromium(VI) removal from solution as a function of time and pH. The initial concentration for all samples was 10 mg/l (as Cr(VI)).

surface charges with decreasing pH, which results in more sorption sites that are available for anions.

After about three weeks reaction time, Cr(VI) equilibrium with the soil was still not achieved. Whether adsorption or a possible reduction of Cr(VI) to Cr(III) was responsible for this slow kinetic uptake behavior was not determined. Bartlett and Kimble (2) found Cr(VI) reduction for their experiments in the presence of soil organic matter and low pH. Both the organic matter content and pH conditions in the present study were such that a reduction reaction was likely to occur. Bloomfield and Pruden (4) showed the strong pH and time dependence

of the reduction process. Their graphical results looked quite similar to Figure I.2. Since the solid phase concentrations for Figure I.2 were calculated from the measured liquid phase concentrations, Cr(VI) uptake due to true adsorption could not be distinguished from the removal due to chromium(VI) reduction to Cr(III) and its subsequent removal.

A similarly strong kinetic behavior is shown in Figure I.3, where Cr(VI) removal is plotted as a function of time and different concentrations. The increased Cr(VI) removal for higher liquid phase concentrations could be explained by increased chromium(VI) reduction to Cr(III). Like in the experiment described

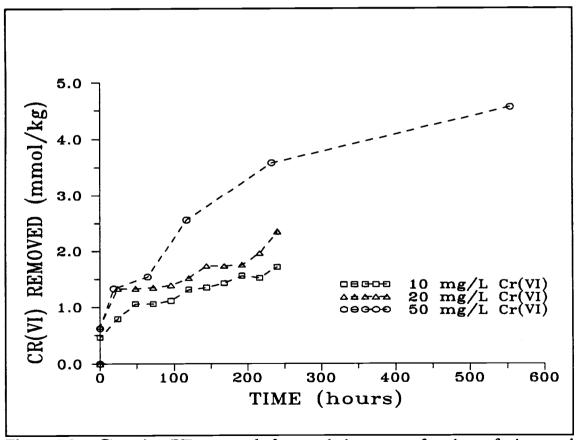


Figure I.3 Chromium(VI) removal from solution as a function of time and concentration at pH 4.5-4.8.

above, it could not be distinguished between Cr(VI) removal due to adsorption and Cr(VI) removal due to reduction.

To further investigate anion sorption kinetics on the test soil, a similar adsorption study for phosphate was conducted using the same soil as in the Cr(VI) experiments. Phosphate was selected because of its similar tetragonal shape to Cr(VI) and its direct competition with chromium(VI) for specific sorption sites (Stollenwerk and Grove (11), Bartlett and Kimble (2)). Additionally, phosphate does not undergo redox reactions and thus removal of phosphate from solution due to redox reactions, which was a suspected cause for kinetically slow Cr(VI) removal, could be excluded. Results of a phosphate adsorption experiment are presented in Figure I.4.

A pH of about 4 was chosen near the original soil pH. The graph indicates a somewhat similar adsorption behavior as observed for chromium(VI). A fast initial uptake was followed by a slower, kinetically controlled uptake. The phosphate removal process, however, in contrast to the Cr(VI) experiments, seemed to be closer to equilibrium after two weeks of reaction time. Under this assumption, about 56% of the phosphate removal occurred fast (within 4 hours), whereas 44% of the removal was kinetically controlled.

To verify that the observed phosphate removal was actually caused by adsorption and to exclude phosphate removal possibly caused by precipitation or complexation, a second set of phosphate batch experiments was conducted using a much lower initial phosphate concentration (0.001 M). For these experiments two sets of soil samples were prepared, one the originally prepared soil described in

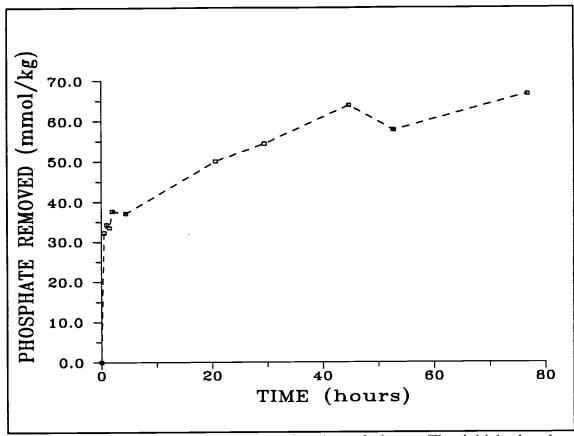


Figure I.4 Phosphate adsorption as a function of time. The initial phosphate concentration was 0.02 M and the pH range was 3.8-4.2.

the Materials section and a second for which the soil was additionally ground in a crucible for 15 minutes to create smaller soil particles. This latter experiment should indicate whether the kinetic behavior of the adsorption reaction was related to chemically controlled or physically controlled reaction kinetics.

A similar pattern of phosphate removal for the experiment utilizing the original soil (0.001 M phosphate) as for the 0.02 M phosphate solution experiment was observed (Figure I.5). The fraction of initial rapid uptake to total uptake was the same for both experiments (0.56). This indicates that no significant precipitation or complexation reactions occurred, since the fractions of initial rapid

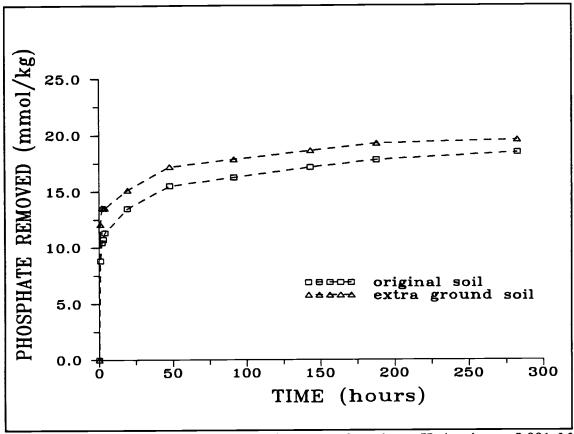


Figure I.5 Phosphate adsorption experiment conducted at pH 4 using a 0.001 M initial concentration. The experiment was performed for two different particle sizes.

uptake and kinetically controlled slow uptake would not remain constant for both experiments. It was concluded that the kinetic behavior in these phosphate experiments was due to adsorption.

This kinetic adsorption behavior could be caused by either a chemical or physical non-equilibrium process (Enfield and Shew (7), Enfield et al. (8)). From Figure I.5 it can be seen that the fraction of initial uptake increased (0.66) as the soil particle size was decreased. The kinetically controlled phase stayed essentially the same. If the kinetic adsorption behavior was caused by a chemical non-equilibrium, e.g. two different kinds of sorption sites, both parts of the

adsorption curve should have increased due to the larger surface area achieved by creating a smaller particle size. It was therefore concluded, that a physical process, likely intraparticle diffusion, was the major cause of the slow kinetic adsorption behavior.

Attempts to measure a chromium(VI) isotherm at true equilibrium failed, since even after 500 hours of reaction time equilibrium was not achieved (Fig. I.2 and Fig. I.3). It was concluded that besides adsorption, chromium(VI) reduction occurred and was the major cause of this non-equilibrium phenomenon. Since the two Cr(VI) removal phenomena could not be differentiated analytically, a Cr(VI) one-hour, non-equilibrium "isotherm" experiment was conducted (Figure I.6). For this experiment it was assumed that adsorption was the major chromium(VI) removal process, since the reduction reaction is a slower, time dependent process. Fitting the data to a Langmuir equation, adsorption maxima and Langmuir coefficients were determined for different pH values (Table I.2). The determined maximum adsorption capacities are most likely underestimated, since the slow kinetically controlled adsorption phase was neglected in the experiment. Using the phosphate batch experiments as a rough estimate, the maximum adsorption capacities for a true equilibrium isotherm for chromium(VI) can be expected to be almost twice the values determined in the one-hour experiments. This estimate would lead to adsorption capacities close to those determined by Griffin et al. (9) for montmorillonite clay.

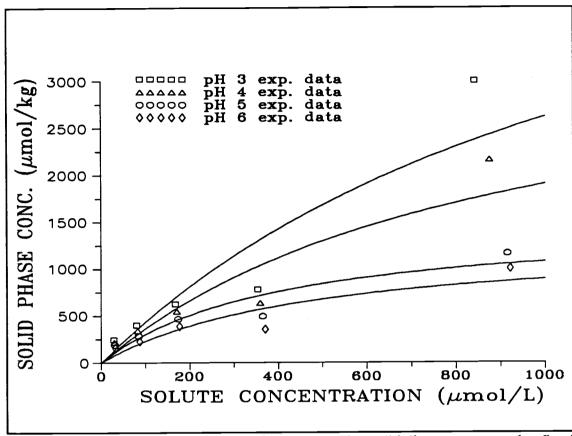


Figure I.6 Chromium(VI) one-hour isotherm. The solid lines represent the fitted Langmuir equations varying pH.

Table I.2 Cr(VI) one-hour Langmuir isotherm results for varying pH.

	pH	3	4	5	6	7	8
$Q_{M}$	(mmol/g)	5.94	3.58	1.49	1.30	1.17	0.84
b	(L/mmol)	0.79	1.13	2.57	2.17	1.54	0.93
Q <sub>м</sub> b	= maximum ads = Langmuir ads						

Soil column experiments were performed for Cr(VI) and Cl. Figure I.7 presents the breakthrough and desorption curves for experiments no. 1, 2 and 4 (Table I.1). The time difference for breakthrough between the chromium(VI) and the Cl tracer experiments indicates the strong adsorption tendency exhibited by chromium(VI). The breakthrough for experiment no. 2 occurred earlier than the breakthrough for experiment no. 1, accompanied by steeper slopes of the breakthrough curve. These effects can be attributed to the higher initial

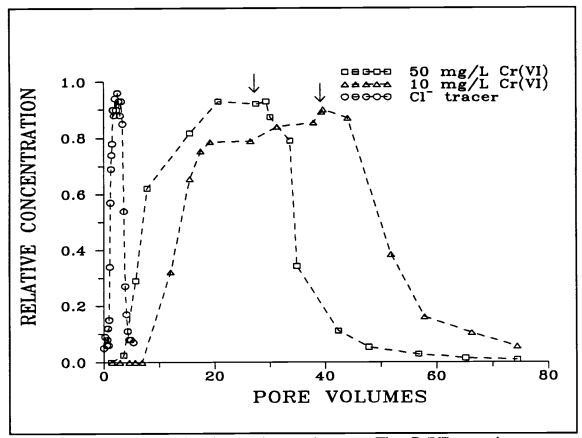


Figure I.7 Soil column breakthrough experiments. The Cr(VI) experiments were performed at pH 3.8-4.2. The desorption was conducted using distilled water. The arrows indicate the beginning of the desorption phases.

chromium(VI) concentration in experiment no. 2. Available sorption sites are more rapidly occupied, thus creating a higher mobility for the Cr(VI) remaining in solution. Noticeable in Figure I.7 is the asymmetry of the chromium(VI) breakthrough curves. The desorption limbs appear to be delayed. This can be attributed to non-equilibrium conditions during the experiment and is referred to as tailing. Similar breakthrough curves for Cr(VI) in alluvium were presented by Stollenwerk and Grove (11).

The breakthrough and desorption curve for experiment no. 3 is presented in Figure I.8. In this experiment, a 0.02 M H<sub>2</sub>PO<sub>4</sub> solution (at pH 4) was used

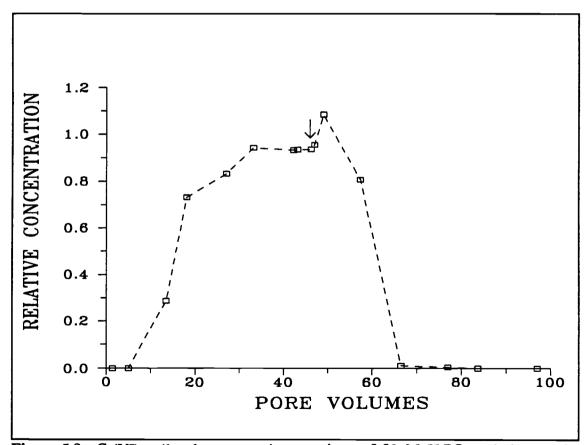


Figure I.8 Cr(VI) soil column experiment using a 0.02 M H<sub>2</sub>PO<sub>4</sub> solution as the extractant. The experiment was conducted at pH 3.8 - 4.2. The arrow indicates the beginning of the desorption phase.

for the chromium(VI) desorption phase instead of distilled water. The phosphate addition caused a Cr(VI) spike in the effluent, exceeding the initial influent Cr(VI) concentration. This spike was followed by a steep decline of Cr(VI) concentration to almost zero. No curve tailing was observed.

Phosphate, as observed in previous studies (Bartlett and Kimble (2), Stollenwerk and Grove (11)), showed a strong tendency to compete with Cr(VI) for adsorption sites. This competition caused Cr(VI) to be desorbed from the soil much more rapidly than observed for the distilled water desorption. The 0.02 M phosphate solution (Experiment no. 3) reduced the chromium(VI) concentration in the effluent to near zero in about 100 pore volumes less than for distilled water (Experiments no. 1 and 2, Table I.3). A mass balance performed on Cr(VI) for these soil column experiments (Table I.3) showed that only slightly more Cr(VI) could be recovered using the phosphate extraction method compared to the distilled water extraction. The amount of Cr(VI) residual in the soil was fairly constant. Relating these column results to the results of the Cr(VI) batch experiments, it seems likely that the unrecoverable amount of chromium(VI) is due to Cr(VI)

Table I.3 Cr(VI) extraction experiments - Mass balances.

Experiment no.	Pore volumes for extraction	Cr(VI) added (mol)	Cr(VI) recovered (mol)	Cr(VI) residual (mol)
1	155.8	1.40*10-4	, 1.20*10 <sup>-4</sup>	2.00*10-5
2	158.3	4.89*10 <sup>-4</sup>	4.77*10 <sup>-4</sup>	1.20*10-5
3	50.8	1.64*10 <sup>-4</sup>	1.50*10 <sup>-4</sup>	1.36*10 <sup>-5</sup>

reduction and is actually present in the soil as Cr(III). The phosphate extraction solution is ineffective for this fraction since Cr(III) at the experimental pH of 4 would be present (and adsorbed onto the soil) in the cationic form. The unrecoverable amount of Cr(VI) could also be due to very tight specific adsorption of Cr(VI). In this case, however, the addition of phosphate should have led to a more significant decrease in the amount of unrecoverable Cr(VI).

## SUMMARY AND CONCLUSIONS

The behavior of Cr(VI) in a natural soil was studied utilizing batch reactor as well as soil column experiments.

Batch experiments were performed to investigate the effects of pH and initial concentration on chromium(VI) removal as a function of time. Difficulties were encountered distinguishing true adsorption from Cr(VI) reduction to Cr(III) and subsequent removal. For this reason, phosphate batch experiments were conducted under the assumption that phosphate exhibits an adsorption behavior comparable to that of Cr(VI). The similar shape of phosphate adsorption observed in previous studies and its competitive sorption effect on Cr(VI) seemed to justify such an assumption.

Soil column experiments were performed for Cr(VI) to more closely simulate field conditions of continuous hydraulic flow through a stationary porous medium. Breakthrough curves were determined for different influent concentrations of Cr(VI) using distilled water as the extracting agent. A chloride tracer breakthrough curve was determined and compared to the Cr(VI) breakthrough results. To investigate the effectiveness of phosphate on Cr(VI) removal from contaminated soil, another Cr(VI) soil column experiment was conducted in which a phosphate solution was used as the extraction agent.

From this study it was concluded that:

- 1. Cr(VI) removal in natural soil suspensions is most likely due to true adsorption as well as Cr(VI) reduction to Cr(III) and subsequent removal.
- 2. Cr(VI) removal increases strongly with decreasing pH and increasing initial Cr(VI) concentration.
- 3. The adsorption process for Cr(VI) likely consists of an initial fast uptake followed by a slower, more time dependent removal. Investigations indicated that this behavior can be attributed to a physical non-equilibrium process.
- 4. Phosphate strongly improved the Cr(VI) removal effectiveness by increasing the Cr(VI) desorption rate significantly, although it was not able to recover significantly more Cr(VI) from a contaminated soil during a soil column experiment.

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### CHAPTER II:

A Kinetic Solute Transport Model For
The Movement Of Chromium(VI) Through Soils

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## **ABSTRACT**

To increase the performance of cleanup methods for chromium contaminated soils, the fate and transport of chromium through soils have to be investigated and better understood. The objective of this work was to develop and verify a solute transport model to predict chromium(VI) movement through soil columns. Of particular interest was the behavior of chromium(VI) in the presence of phosphate, another adsorbing anion. The model was developed under the assumption of one-dimensional, steady-state saturated groundwater flow through a homogeneous porous medium. Sorption of chromium(VI) and phosphate are described with a physical non-equilibrium model (mobile and immobile water phases) and a competitive Langmuir isotherm. Numerical solutions of the transport

equations were obtained by the partially implicit finite difference method.

Computer simulations were fitted to experimental breakthrough curves using estimates for model parameters which could not be determined independently in experiments. Chromium(VI) breakthrough experiments were successfully modeled for varying chromium concentrations as well as for the case of phosphate competition. This study demonstrated that the use of phosphate as a desorption agent is a possible way of increasing chromium(VI) desorption rates and, consequently, the effectiveness of chromium removal from contaminated soils.

Additional Word Index: chromium, solute transport, physical non-equilibrium, adsorption, competitive Langmuir equation.

### INTRODUCTION

Groundwater is an important freshwater resource. Over the last decade, more and more cases of groundwater contamination by heavy metals, in particular chromium, have been discovered. Such contaminations pose a severe threat to human health as well as to the environment. To predict the movement of chromium through soils and aquifer systems, its transport mechanisms and particularly its interactions with the soil matrix have to be investigated and better understood. The objective of this work was to develop a solute transport model to describe the movement of chromium through laboratory soil columns, and to verify the model using experimentally determined breakthrough curves (BTC's). Of specific interest was the modeling of chromium(VI) solute transport in the presence of phosphate. Experimental results of an ongoing feasibility study on chromium extraction from a contaminated aquifer indicated that a higher Cr(VI) removal efficiency is possible if phosphate is added to the leaching water (Nelson (19)).

A variety of solute transport models for both organic and inorganic species have been proposed. Most of these models incorporate the same dispersion-advection transport processes. They differ, however, in the way that sorption / desorption reactions are modeled. A review of sorption models for reactive solutes in soil was given by Travis and Etnier (24). Both equilibrium and first-order kinetic models were discussed. Equilibrium models were originally developed to describe sorption isotherms for gases and organic solutes. Bar-Yosef

(14) derived a sorption model based on a competitive Langmuir isotherm for ionic species and applied it successfully to pH dependent zinc adsorption by soils. This model did not include the effect of the electric potential (Ψ) and the electrolyte concentration on adsorption. Bar-Yosef believed that the effect of Ψ was to some extent accounted for by the binding constants of the different ions. A transport model based on ion exchange was presented by Valocchi et al. (25). Sorption processes for ionic species were modeled employing the well known principle of ion exchange selectivity. Although successful for the cases considered, the authors realized that the ion selectivity coefficients are typically not constant but instead are a function of the sorbed phase concentration. These coefficients are also highly variable in natural soil systems, thus hard to estimate. Cederberg et al. (15) presented a solute transport model that incorporated ion exchange and surface complexation coupled with a chemical equilibrium model. Good agreement was found between predicted and measured concentrations of cadmium, chloride and bromide in laboratory soil columns.

For many transport problems, the assumption of equilibrium sorption is not valid (Van Genuchten (29)). Two classes of non-equilibrium models have been developed: chemical non-equilibrium and physical non-equilibrium.

Derivations of these models are in Van Genuchten and Cleary (28) and Van Genuchten (29).

The chemical non-equilibrium model assumes that there are two types of sorption sites. For one type of site the sorption reaction is assumed to be a slow, kinetically controlled process. For the other type of site sorption is assumed to be

a kinetically faster or even an instantaneous equilibrium process. Such a two - site model was recently used by Selim and Amacher (23) to model chromium(VI) transport through three different soils. Kinetic Langmuir equations were used to describe the sorption reactions. This approach was partially successful in describing data from miscible displacement experiments. Grove and Stollenwerk (17) modeled chromium(VI) movement through alluvial materials coated with iron-oxide and hydroxide. They used a one-site chemical non-equilibrium model based on a kinetic Langmuir equation. This model agreed better with experimental data than when sorption was assumed to be at equilibrium.

Physical non-equilibrium models usually divide the total water content of a soil into a mobile and an immobile region and assume that the kinetic sorption behavior is caused by a physical mechanism, typically a diffusional process across the immobile water layer to the solid surface. Wu and Gschwend (30) recently presented evidence that the sorption kinetics of hydrophobic organic compounds on natural sediments are controlled by such an intraparticle diffusion.

These two major groups of non-equilibrium solute transport models were compared by Nkedi-Kizza et al. (20). When a linear sorption reaction was used, the analytical solutions for the two models are equivalent. These authors concluded that the difference between the two models is of little practical importance and suggest that the occurrence of specific processes can only be verified with microscopic measurements.

#### THEORY

# Model Development

Based on the results of batch experiments (Schroth et al. (22)) it was believed that the kinetic sorption behavior of chromium(VI) is controlled by an intraparticle diffusional process. Therefore a physical non-equilibrium solute transport model was developed under the assumption of one-dimensional, steady-state saturated groundwater flow through homogeneous porous media. A derivation of a similar model is in Van Genuchten (29).

The model assumes that the total water content of the soil is divided into a mobile phase and an immobile phase, so that

$$\Theta = \Theta_{\rm m} + \Theta_{\rm im} \tag{1}$$

where  $\Theta$  = total volumetric water content (cm<sup>3</sup>/cm<sup>3</sup>)

 $\Theta_{\rm m}$  = volumetric water content of the mobile phase (cm<sup>3</sup>/cm<sup>3</sup>)

 $\Theta_{im}$  = volumetric water content of the immobile phase (cm<sup>3</sup>/cm<sup>3</sup>)

where the subscripts "m" and "im" identify the mobile and immobile phases, respectively. Conceptually, the mobile pore water content is understood to include the water in the pore space between soil particles, whereas the immobile water

content includes the water in the pore space within particles.

Solute transport due to advection and dispersion is limited to the mobile water phase. The following dispersive-advective transport equation is well documented in the literature (Van Genuchten and Wierenga (27), De Smedt and Wierenga (16), Rubin (21)) and has found wide application:

$$\frac{\partial c_m}{\partial t} = D \frac{\partial^2 c_m}{\partial x^2} - v \frac{\partial c_m}{\partial x}$$
 (2)

where  $c_m = \text{solute liquid phase concentration (mg/l)}$ 

t = time(h)

D = dispersion coefficient (cm<sup>2</sup>/h)

x = distance (cm)

v = average pore water velocity (cm/h)

The solute is transferred between the mobile and immobile water phases by an intraparticle diffusional process. In the mobile water phase this process is assumed to be controlled by the first-order expression:

$$\frac{\partial c_{m}}{\partial t} = -\frac{\alpha}{\Theta_{m}} (c_{m} - c_{im})$$
 (3)

and in the immobile water phase by:

$$\frac{\partial c_{im}}{\partial t} = \frac{\alpha}{\Theta_{im}} (c_m - c_{im})$$
 (4)

where  $\alpha$  = rate coefficient for diffusion across the immobile water layer (h<sup>-1</sup>)

The rate coefficient is assumed to be the same for all solutes, thus neglecting the effect of molecular radius of solute species on the diffusional process.

Sorption occurs in both the mobile and immobile water phases. It is assumed that a certain mass fraction of the soil, f, is accessible in the mobile water phase and a fraction, (1-f), is accessible in the immobile phase, so that

$$S = fS_{m} + (1-f)S_{im}$$
 (5)

where S = total adsorbed concentration (mg solute/g soil)

S<sub>m</sub> = adsorbed concentration for solid fraction which is in contact with mobile water phase (mg solute/g soil)

S<sub>im</sub> = adsorbed concentration for solid fraction which is in contact with immobile water phase (mg solute/g soil)

The sorption reaction for both phases is assumed to be at equilibrium. This is referred to as the local equilibrium assumption (Jennings and Kirkner (18), Valocchi (26)). The change of solute concentration due to sorption is

$$\frac{\partial c_m}{\partial t} = -f \frac{\rho}{\Theta_m} \frac{\partial S_m}{\partial t}$$
 (6)

in the mobile phase and

$$\frac{\partial c_{im}}{\partial t} = - (1-f) \frac{\rho}{\Theta_{im}} \frac{\partial S_{im}}{\partial t}$$
 (7)

in the immobile phase,

where  $\rho$  = bulk density of soil (g soil/cm<sup>3</sup>)

The governing partial differential equation for the mobile phase can be obtained by combining Eq.(2), Eq.(3) and Eq.(6):

$$\frac{\partial c_{m}}{\partial t} = D \frac{\partial^{2} c_{m}}{\partial x^{2}} - v \frac{\partial c_{m}}{\partial x} - f \frac{\rho}{\Theta_{m}} \frac{\partial S_{m}}{\partial t} - \frac{\alpha}{\Theta_{m}} (c_{m} - c_{im})$$
 (8)

and for the immobile water phase by combining Eq.(4) and Eq.(7):

$$\frac{\partial c_{im}}{\partial t} = -(1-f) \frac{\rho}{\Theta_{im}} \frac{\partial S_{im}}{\partial t} + \frac{\alpha}{\Theta_{im}} (c_m - c_{im})$$
 (9)

To solve Eq.(8) and Eq.(9), a sorption relationship has to be specified. For this solute transport model a competitive Langmuir equation, as presented by Bar-Yosef (14), was chosen to represent the sorption reaction. This type of equation, unlike the frequently chosen linear isotherm, allows for competition between different solutes for the available sorption sites. When applied to the mobile-immobile water phase model, the sorption equations become :

$$S_{m}(i) = \frac{Q_{M} b(i) c_{m}(i)^{1/z(i)}}{1 + \sum b(i) c_{m}(i)^{1/z(i)}}$$
(10)

and

$$S_{im}(i) = \frac{Q_{M} b(i) c_{im}(i)^{1/z(i)}}{1 + \sum b(i) c_{im}(i)^{1/z(i)}}$$
(11)

where  $Q_M = \text{maximum adsorption capacity (} \mu g \text{ solute/} g \text{ soil)}$ 

b = Langmuir adsorption constant (1/mg)<sup>2</sup>

z = valence of a species

i = solute species

The complete system is now defined by a set of four equations (8, 9, 10, 11) for each solute species considered.

## Numerical Solution

Equations (8), (9), (10) and (11) form a system of coupled non-linear equations for each solute species. It is impossible to obtain exact analytical solutions for such a system (Van Genuchten and Cleary (28)). Therefore numerical methods have to be used.

A numerical solution approximation was developed using a partially implicit finite difference method, where the soil column is imagined to be divided into a certain number of adjacent cells. For each of these cells, dispersion,

advection and diffusional mass transfer between the mobile and immobile water phase is first computed explicitly, using equilibrium solution concentrations computed in the previous time step. Then the sorption reaction is accounted for by equilibrating the resulting solution concentrations with the solid phase using the Newton-Raphson method.

The finite difference approximation for the mobile water phase is given by:

$$c(i,n)_{m}^{t+\Delta t} = c(i,n)_{m}^{t} + \frac{D\Delta t}{\Delta x^{2}} \left[ c(i,n-1)_{m}^{t} - 2c(i,n)_{m}^{t} + c(i,n+1)_{m}^{t} \right]$$

$$+ \frac{v\Delta t}{\Delta x} \left[ c(i,n-1)_{m}^{t} - c(i,n)_{m}^{t} \right] - \frac{f\rho}{\Theta_{m}} \left[ S(i,n)_{m}^{t+\Delta t} - S(i,n)_{m}^{t} \right]$$

$$- \frac{\alpha \Delta t}{\Theta_{m}} \left[ c(i,n)_{m}^{t} - c(i,n)_{im}^{t} \right]$$
(12)

and for the immobile water phase by:

$$c(i,n)_{im}^{t+\Delta t} = c(i,n)_{im}^{t} - (1-f)\frac{\rho}{\Theta_{im}}[S(i,n)_{im}^{t+\Delta t} - S(i,n)_{im}^{t}] + \frac{\alpha \Delta t}{\Theta_{im}}[c(i,n)_{m}^{t} - c(i,n)_{im}^{t}]$$
(13)

where n = cell number

 $\Delta x = length of one cell (cm)$ 

t = current calculation time (h)

 $\Delta t = length of time step (h)$ 

The initial condition for both water phases is

$$c_m(x,t=0) = c_{im}(x,t=0) = c_i$$
 (14)

where  $c_i$  = initial concentration inside the soil column (mg/l)

Boundary conditions apply only to equation (12), since only the mobile water phase is assumed to be in direct contact with the boundaries. A third-type, constant flux boundary condition was used to model the upper boundary:

$$(-D\frac{\partial c}{\partial x} + vc)_{x=0} = vc_0$$
 (15)

where  $c_0 = \text{column influent concentration (mg/l)}$ 

The lower boundary was modeled using

$$\frac{\partial c}{\partial x}(L,t) = 0 \tag{16}$$

where L = column length (cm)

These boundary conditions and their influence on the solution of equations (12) and (13) were discussed by Van Genuchten and Wierenga (27) and Van Genuchten (29). These authors recommended the above described boundary conditions particularly for the size of soil columns used in this study.

#### EXPERIMENTAL APPROACH

Continuous flow soil column experiments and batch reactor experiments were performed using soil samples taken from an area close to a chromium contamination site to assure that the uncontaminated soil was similar to the contaminated soil of the site in terms of soil classification and soil properties.

Column experiments were performed for chromium(VI) to investigate its transport behavior with and without phosphate present during the desorption. To determine the dispersivity in the soil columns, a tracer breakthrough experiment was conducted using chloride (Cl') as the tracer ion. The specific conditions, under which these column experiments were conducted, are listed in Table II.1. Batch experiments were conducted for both chromium(VI) and phosphate to determine sorption parameters.

Experimental procedures and analytical methods for these experiments were discussed in detail elsewhere (Schroth et al. (22)).

Table II.1 Parameters for continuous flow soil column experiments

		Loading time (mob. pore volumes)	
1	14.6	2.58	H <sub>2</sub> O
r(VI) column	experiments:		
Experiment	Influent Cr(V	/I) Loading time (mob. pore volumes)	
Experiment	Influent Cr(V		

#### RESULTS AND DISCUSSION

# Tracer Breakthrough Experiment

The experimental data of experiment no. 1 were imported into a computer program presented by Van Genuchten (29) to determine the dispersivity using a non-linear, least-square analysis. The transport equation to which the data was fitted contained a linear equilibrium sorption term. Since a tracer is assumed not to adsorb onto a solid phase, the retardation factor for the program input was set to one. The program output provided a dispersion coefficient from which a dispersivity of 0.30 cm was computed. The output also provided a retardation factor of 1.18, indicating that the tracer did not behave ideally, but showed a small tendency to adsorb to the soil. Such a retardation could be caused by a non-specific ion exchange process.

The shape of the experimental chloride breakthrough curve is of particular interest to the mobile-immobile water phases model (Figure II.1). The symmetric shape of the curve indicates little kinetic behavior for Cl movement. According to the mobile-immobile water phase model, the Cl anions are assumed to undergo the diffusional process between the mobile and immobile water phases. Since Cl exhibited only a small sorption tendency, equilibrium between the two phases is achieved more rapidly than for anions which adsorb more strongly onto the soil. However, if the water content of the immobile phase was large compared to the

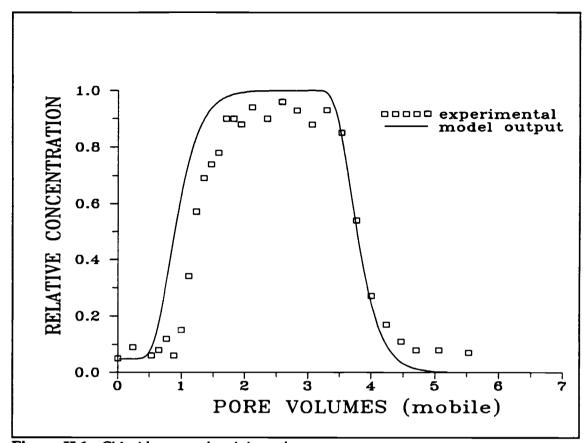


Figure II.1 Chloride tracer breakthrough curve.

water content of the mobile phase, the diffusional mass transfer into the immobile phase would require more time and therefore introduce more kinetics to the solute transport. This would create a larger asymmetry of the breakthrough curve, leading to more curve tailing. The little tailing observed therefore suggests a small water content of the immobile water phase compared to the water content of the mobile phase.

The output from the finite difference model is also shown in Figure II.1.

During the computation, it was assumed that the tracer moving through the soil

column would not interact with the solid phase. This caused the breakthrough to

appear earlier than experimentally observed. Experimental data and model output showed a better fit for the desorption part of the breakthrough curve. The discrepancy between observed and predicted concentration in the two parts of the breakthrough curve could be due to chemical hysteresis (Van Genuchten and Cleary (28)). Chemical hysteresis is caused by a difference between adsorption and desorption rates of a reacting solute. An experimentally observed background level of Cl<sup>-</sup> is the cause for the deviation between model and experiment at the end of the desorption limb of the breakthrough curve. The Cl<sup>-</sup> background level might be caused by mineral soil particles dissolving into solution during the solute transport process.

# Chromium(VI) Breakthrough Curves

Chromium(VI) anions are known to adsorb strongly onto soil (Grove and Stollenwerk (17), Selim and Amacher (23)). Batch sorption experiments exhibited a strong kinetic behavior, showing fast initial Cr(VI) uptake from solution followed by a slower time dependent uptake (Schroth et al. (22)). This kinetic behavior is assumed to be controlled by a rate coefficient ( $\alpha$ ) for diffusional mass transfer across an immobile water phase. To demonstrate the effect of  $\alpha$  on the solute transport of reactive solutes, chromate breakthrough curves were simulated for varying values of  $\alpha$  using the numerical model. The breakthrough curve for  $\alpha$ =0.0 represents equilibrium conditions (Figure II.2). Although no kinetics are included,

the BTC is asymmetric. This is due to the non-linear sorption relationship used and was explained by Van Genuchten and Cleary (28). When kinetic behavior is

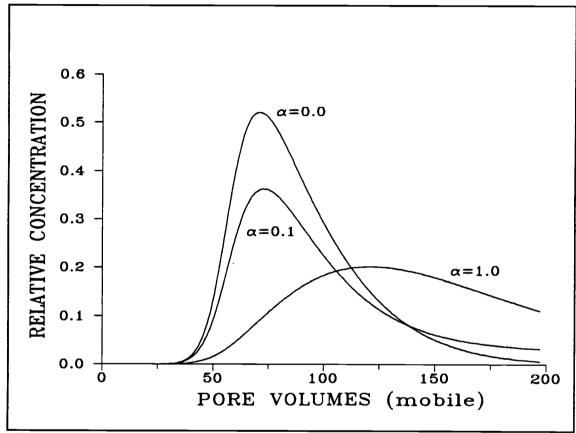


Figure II.2 Chromium(VI) breakthrough curve simulations for different values of the rate coefficient  $\alpha$ .

included ( $\alpha > 0$ ), the breakthrough curves show an increasingly asymmetric shape. Unfortunately, no laboratory experiment was found appropriate to measure this rate coefficient independently. In previously conducted studies,  $\alpha$  was either estimated (Selim and Amacher (23)) or fitted (Wu and Gschwend (30), Van Genuchten (29)).

Chromium(VI) breakthrough curves were modeled using parameters previously determined in laboratory batch experiments (Schroth et al. (22)). Two chromate sorption parameters, the maximum sorption capacity and the Langmuir adsorption constant, were determined in an one-hour batch isotherm experiment. Efforts to measure these parameters using a long-term, true equilibrium isotherm failed because even after a reaction period of three weeks equilibrium could not be achieved. Therefore it had to be assumed that only a fraction of chromium was adsorbed during the one-hour experiment, and thus the maximum adsorption capacity was underestimated in this experiment. Comparable phosphate isotherm experiments showed that the maximum adsorption capacity was almost twice as high in a 14 day long-term experiment than in an one-hour experiment. For these phosphate experiments it could be assumed that sorption was the only major mechanism removing the phosphate from solution (Schroth et al. (22)).

Another important parameter for the model was the mass fraction of solids that is in contact with the mobile water phase (f). In a previous study Selim and Amacher (23) assumed this fraction to be the same as the fraction of pore water content in the mobile phase to the total water content of the soil. For this study the fraction, f, was estimated to be 0.56 using phosphate batch sorption experiments (Schroth et al. (22)). This value is smaller than the ones proposed by Selim and Amacher (23), meaning that a smaller portion of the soil is available for instantaneous equilibrium sorption. The fraction of the mobile water content to the total water content,  $(\Theta_m/\Theta)$ , as a result of the tracer experiment, was set to 0.95, larger than the ones determined by Selim and Amacher (23).

The parameters which were determined in the laboratory experiments, in particular the sorption parameters, are conditional by nature and therefore have to be treated cautiously when applied to the solute transport model. The sorption parameters, as presented, are highly pH dependent. The solute transport model, however, does not account for a change in sorption parameters due to changing pH conditions. The model could therefore lead to errors in case of large pH changes during a soil column experiment. For all the experiments presented in this paper, pH changes were small enough to be neglected.

The parameters presented in Table II.2 were used to model a chromium(VI) breakthrough experiment previously conducted in laboratory experiment no.2 (Table II.1). Figure II.3 shows the strong sorption tendency of chromium(VI), requiring more than 20 mobile pore volumes to reach a relative effluent concentration close to 1. Also visible is the tailing of the breakthrough

Table II.2 Solute transport model parameters.

arame	ter 	Cr(VI) experiments	Cl <sup>-</sup> experiment 
d	(cm)	0.30	0.30
ρ	$(g/cm^3)$	1.09	1.09
f		0.56	0.56
$\Theta_{m}/\Theta$		0.95	0.95
α	(hour-1)	0.20	0.20
$Q_{M}$	(µmol/g)	5.00	0
b <sub>Cr</sub>	(l/mmol)	1.13	
$b_{Ph}$	(l/mmol)	1.24	
z		1	1

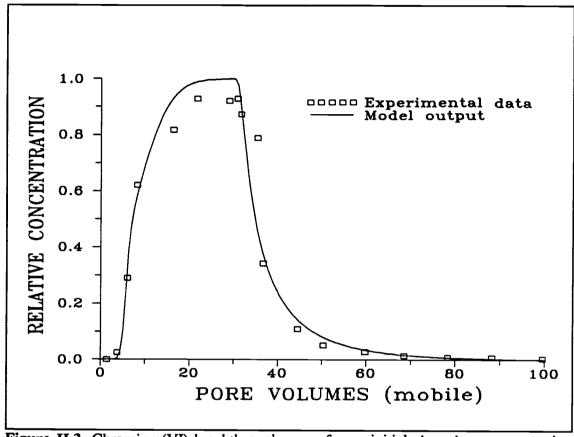


Figure II.3 Chromium(VI) breakthrough curve for an initial chromium concentration of 50 mg/l (as Cr) and desorption using distilled water.

curve, which to a large degree is due to the non-equilibrium condition inside the column and the strong sorption tendency of Cr(VI). Tailing occurs because the release of Cr(VI) adsorbed to surface sites which are in contact with the immobile water phase into the mobile water phase is slowed by the diffusional mass transfer process between the two phases, requiring more time for this Cr(VI) fraction to reach the mobile water phase. That causes the BTC not to decrease as fast as it would in an equilibrium process. This effect was described in detail by Van Genuchten and Cleary (28). The mass transfer coefficient  $\alpha$  for the model was set to 0.2 hour by adjusting the slope of the modeled breakthrough curve to the

experimental data. Using this  $\alpha$  value, the model output agreed best with the experimental results. A deviation between model and experiment can be observed in the region where the relative concentration is close to 1. Here the model predicted a relative concentration of 1, whereas the experimental effluent chromium(VI) concentration never reached the influent concentration throughout the experiment. Assuming little experimental uncertainty, this could indicate that there is another chromium(VI) sink not accounted for in the model. Mass balances performed on the experimental data supported this suggestion. In all chromium(VI) experiments performed, a small percentage of chromate could never be recovered. This could be due to chromium(VI) reduction and is discussed elsewhere (Schroth et al. (22)).

Experiment no.3 (Table II.1) was modeled using the same parameters (Table II.2) as were used for experiment no.2. The model output agrees well with the experimental data (Figure II.4), although the breakthrough in the second experiment occurs later than in the first due to the different initial chromium(VI) concentration.

# Chromium(VI) Breakthrough Experiment Including Anion Competition

To investigate the influence of a competitive anion on the desorption of chromium(VI) from soil, chromium(VI) breakthrough experiment no.4 (Table II.1) was conducted in which a 0.02 M phosphate solution, instead of distilled water,

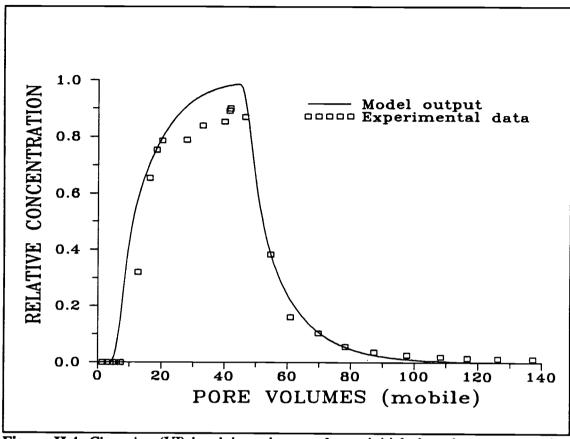


Figure II.4 Chromium(VI) breakthrough curve for an initial chromium concentration of 10 mg/l (as Cr) and desorption using distilled water.

was injected to desorb the chromate from the soil beginning at 48.5 mobile pore volumes. The results are presented in Figure II.5. The experimental data show a chromate concentration spike after 49 mobile pore volumes followed by a rapid decrease in concentration. This effect can be attributed to competitive sorption by the phosphate anion. The solid line in Figure II.5 shows the model prediction for the breakthrough using the phosphate solution for desorption. Similarly to the experiment, a chromate spike occurs at about 49 mobile pore volumes followed by a fast decline of chromate in the effluent to almost zero at 57 mobile pore volumes. The good fit between experiment and model demonstrated that the

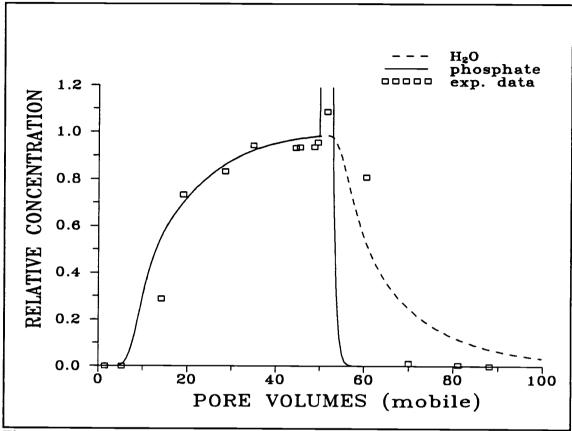


Figure II.5 Chromium(VI) breakthrough curve for an initial concentration of 10 mg/l (as Cr). In the experiment phosphate was injected for desorption. The solid/dashed lines show the model output for phosphate and H<sub>2</sub>O desorption.

competitive Langmuir equation successfully modeled the anion competition for the available sorption sites. The dotted line shows the breakthrough prediction for chromate when distilled water is used for desorption (see also Fig. II.4) and indicates a much slower process compared to the one with phosphate competition.

The competitive Langmuir equation, although used successfully in the this case, has certain limitations. There is only one maximum sorption capacity  $Q_M$  used in the Langmuir term, thus neglecting soil selectivity effects for different anions. In the case above, the estimated maximum sorption capacity of Cr(VI) at pH 4 (Schroth et al. (22)) was used in the model equations for both Cr(VI) and

phosphate computations. Phosphate, however, adsorbs distinctively stronger onto the soil used in the experiments than Cr(VI). Therefore the phosphate liquid phase concentration was overestimated in the competitive model, causing an overestimation of the phosphate competition effect on chromium(VI) adsorption. Attempts to model the Cr(VI) breakthrough using the maximum sorption capacity determined for phosphate failed even after adjusting the valence z for Cr(VI) adsorption empirically to simulate a higher site coverage per anion adsorbed.

The successful use of the competitive Langmuir equation for the case above was due to the high phosphate influent concentration during the Cr(VI) extraction process. Although the true phosphate liquid phase concentration was smaller than assumed in the model calculations, it was still large compared to the Cr(VI) liquid phase concentration, therefore competing strongly for available adsorption sites.

### SUMMARY AND CONCLUSIONS

A physical non-equilibrium solute transport model was developed to predict the movement of chromium(VI) through soil columns. The model incorporates mobile and immobile water phases to account for dispersion, advection, sorption and diffusional mass transfer between the two water phases. The sorption process was modeled under a local equilibrium assumption using a competitive Langmuir equation. To verify the model, a finite difference approximate solution to the nonlinear system of equations was developed to predict breakthrough curves. These BTC's were compared to experimentally determined soil column BTC's.

From this study the following was concluded:

- 1. A physical non-equilibrium solute transport model can be used to successfully model Cr(VI) soil column breakthrough experiments. The use of a non-equilibrium model for modeling Cr(VI) breakthrough experiments is essential.
- A local equilibrium assumption can be used to model adsorption as an equilibrium process under the conditions present during the soil column experiments.

- 3. The competitive Langmuir equation was used successfully to model adsorption during pure chromium(VI) solute transport as well as competitive adsorption during Cr(VI) transport in the presence of phosphate.
- 4. The study demonstrated that the use of phosphate as a desorption agent is a possible way of increasing Cr(VI) desorption rates and, consequently, the effectiveness of chromium removal from contaminated soils.

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### **NOTATION**

The following symbols are used in this paper:

b = Langmuir adsorption constant

c = solute liquid phase concentration

c<sub>i</sub> = initial concentration inside the soil column

 $c_0$  = column influent concentration

d = soil column dispersivity

D = dispersion coefficient

f = mass fraction of soil which is accessible in the mobile water phase

L = soil column length

 $Q_{M}$  = maximum adsorption capacity

S = adsorbed concentration

t = time

v = average pore water velocity

x = distance in flow direction

z = valence of a species

 $\alpha$  = rate coefficient for diffusion across the immobile water layer

 $\Delta t$  = finite difference time increment

 $\Delta x$  = finite difference distance increment

 $\Theta$  = volumetric water content

 $\rho$  = bulk density of soil

# Subscripts

i = solute species

n = finite difference cell number

m = symbol for the mobile water phase

im = symbol for the immobile water phase

# Superscripts

t = time

 $\Delta t$  = finite difference time increment

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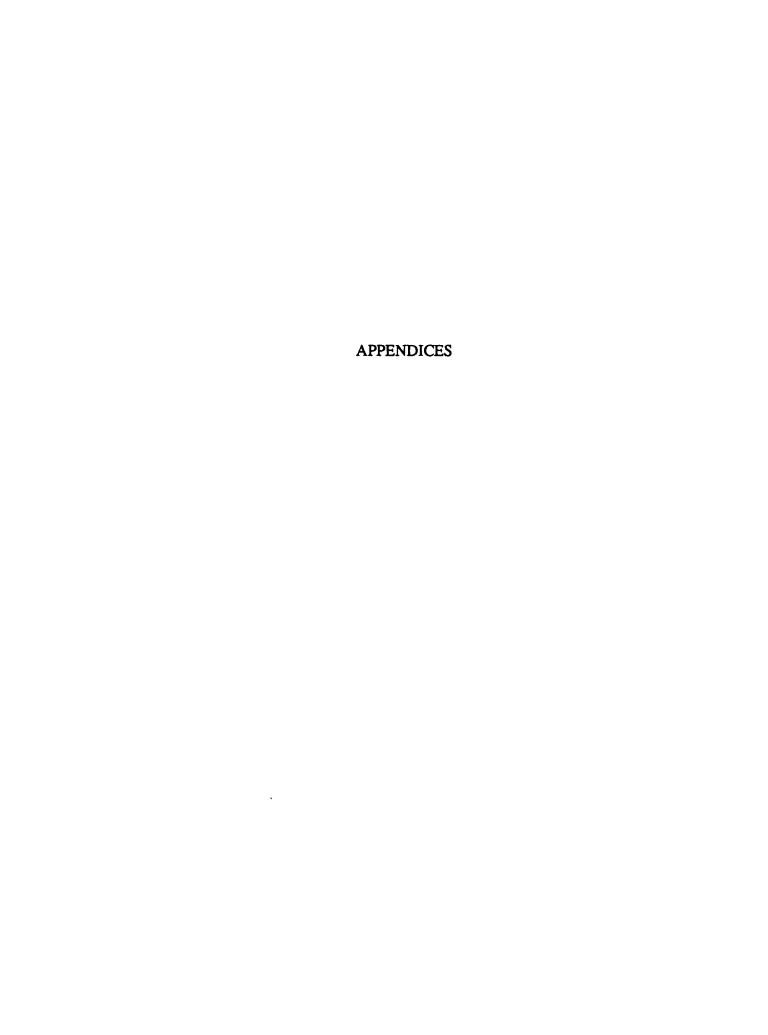
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# APPENDIX A

# LABORATORY RESULTS

This appendix contains tables with the experimental data of the batch reactor and soil column experiments for Cr(VI), phosphate and chloride.

0.17

4.24

Table A.1 Chloride tracer experiment — Experimental data.

Chloride tracer test flowrate : 11 ml/h
pore vol. : 24.6 ml column length: 5.5 cm diameter: 3.2 cm influent C = 14.6 mg/l (as NaCl) pV = 24.6 (100%) pV=23.37 (95%)cum. pore rel. cum. pore rel. sample time time cum. pore Conc average volume mg/L (h) conc. volume volume conc. 0.00 0.71 0.23 1.32 0.51 0.88 0.00 0.00 0.00 0.05 0.00 1.00 0.50 0.23 0.09 0.24 1 2 1.25 1.13 0.51 0.06 0.53 1.50 1.38 0.62 1.20 0.08 0.65 0.62 0.62 1.20 0.73 1.74 1.75 1.63 0.73 0.76 0.12 0.85 0.92 5 2.00 1.88 0.85 0.06 0.88

0.09 0.06 0.08 0.12 0.06 2.25 2.13 0.96 2.17 5.03 6 0.96 0.15 1.00 0.15 1.07 1.07 7 2.50 2.38 0.34 1.12 0.34 8 2.75 2.63 1.19 8.34 1.19 0.57 1.24 0.57 9 3.00 1.30 10.03 1.30 0.69 2.88 1.35 0.69 10 3.25 3.13 1.41 10.85 1.41 0.74 1.47 0.74 11 3.50 3.38 1.53 11.38 1.53 0.78 1.59 0.78 3.63 12 3.75 1.64 13.07 1.64 0.90 1.71 0.90 13 4.00 3.88 1.75 13.18 1.75 0.90 0.90 1.82 14 4.25 4.13 1.87 12.91 1.87 0.88 1.94 0.88 13.75 15 4.75 4.50 2.04 2.04 0.94 2.12 0.94 16 5.25 5.00 2.26 13.15 2.26 0.90 0.90 2.35 5.50 17 5.75 2.49 13.95 2.49 0.96 2.59 0.96 6.00 6.25 2.71 13.59 2.71 0.93 2.82 0.93 6.75 6.50 2.94 12.80 2.94 0.88 3.06 0.88 3.17 3.39 7.25 7.00 3.17 13.53 0.93 3.29 0.93 12.41 7.75 7.50 3.39 3.53 0.85 0.85 7.94 8.25 8.00 3.62 3.62 0.54 3.77 0.54 8.75 8.50 3.84 3.99 4.00 3.84 0.27 0.27

2.51

4.07

0.17

\* desorption using distilled water

9.00

9.25

4.07

Table A.2 Phosphate batch experiment 1 — Experimental data.

Phosphate batch test no.1 add: 1 g soil 25 ml liquid at pH = 4 Initial Conc.: 1743 mg/L as H2PO4-(about 0.02 M) Phosphate Solid Solid Conc. time sample Conc. Conc. (h) (mg/g) (mmo1/kg) no. (mg/1)1743 0.0 1618 3.1 1610 3.3 1613 3.3 0 0 0.00 0.50 1.00 1.50 2.00 4.42 32.2. 34.29 1 2 1597 3.7 3.6 37.64 4 5 1599 37.12 20.58 1549 4.9 50.02 6 29.42 44.75 1532 5.3 54.40 8 1496 6.2 63.81 1519 52.75 57.75 9 5.6 76.75 10 1485 6.5 66.64

Table A.3 Phosphate batch experiments 2 and 3 — Experimental data.

Phosphate batch tests II and III 25 ml of solution w/ initial conc.: 100 mg/L as H2PO4-1 g of soil (fine = Fi ; coarse = Co) liquid liquid solid solid sample time conc.(Co) conc.(Fi) conc.(Co) conc.(Fi) no. (h) (mg/1) (mg/1) (mmol/kg) (mmol/kg) 100 100 8.83 10.45 10.73 65.76 1.00 53.11 12.09 2.00 59.46 47.28 13.59 58.38 3.00 47.61 13.51 4.00 56.09 47.50 11.32 13.54 5 19.25 47.66 15.13 41.31 13.49 47.75 39.92 33.33 15.49 6 17.19 91.25 36.99 30.76 16.24 17.85 8 143.25 33.52 27.76 17.14 18.62 9 187.92 31.08 25.28 17.77 19.26 282.50 10 28.48 24.16 18.44 19.55

Table A.4 Chromium(VI) batch experiment 1 — Experimental data.

Sorption Rate at Various pH's: Compiled from q vs. pH curves at 6 rxn. times 4 hrs-3 weeks rxn. All rxns. begin with 25 mls\*10 ppm Cr(VI) in wellwater Maximum surface density= 4.81E-01 cmol/kg soil Surface Surface Rxn time, Density, Density, [Cr]eq [Cr]eq [Cr], рΗ hours cmol/kg umol/kg mo1/1 ppm umo1/1: -------------------0.25 9.13 1.76E-04 175.5224 3 0.042 420 3 1 0.062 624 8.70 1.67E-04 167.3624 3 4 0.085 850 8.23 1.58E-04 158.3224 24 1.19E-04 0.184 1840 6.17 118.7224 3 45 0.190 1900 6.05 1.16E-04 116.3224 3 166 0.364 3640 2.43 4.67E-05 46.72248 3 0.450 0.64 1.23E-05 331 4500 12.32248 3 500 0.472 4720 0.18 3.52E-06 3.522486 4 0.25 0.0415 415 9.14 1.76E-04 175.7224 1 4 0.0545 545 8.87 1.71E-04 170.5224 4 4 0.064 640 8.67 1.67E-04 166.7224 24 4 0.118 1180 7.55 1.45E-04 145.1224 45 0.135 1350 7.19 1.38E-04 138.3224 4 1.06E-04 5.53 106.3224 166 0.215 2150 4 331 0.260 2600 4.59 8.83E-05 88.32248 4 500 0.310 3100 3.55 6.83E-05 68.32248 5 0.25 0.04 400 9.17 1.76E-04 176.3224 5 0.0466 466 9.03 1.74E-04 173.6824 1 5 4 0.050 500 8.96 1.72E-04 172.3224 5 24 0.086 8.21 860 1.58E-04 157.9224 5 45 0.100 1000 7.92 1.52E-04 152.3224 5 166 0.139 1390 7.11 1.37E-04 136.7224 5 331 0.180 1800 6.26 1.20E-04 120.3224 5 500 0.225 2250 5.32 1.02E-04 102.3224 6 0.25 0.035 350 9.27 1.78E-04 178.3224 6 1 0.0387 387 9.20 1.77E-04 176.8424 6 0.040 400 9.17 1.76E-04 176.3224 4 6 24 0.058 580 8.79 1.69E-04 169.1224 6 45 0.072 720 8.50 1.64E-04 163.5224 6 166 0.08 800 8.34 1.60E-04 160.3224

7.50

1.44E-04

6.98 1.34E-04

144.3224

134.3224

6

6

331

500

0.120

0.145

1200

1450

Table A.5 Chromium(VI) batch experiment 2 — Experimental data.

Reaction time/h		Final	qe, mmol/q	
0.00	250	250	0.000	
0.02	250	226	0.462	
24.00	250	209	0.789	
48.00	250	195	1.058	
72.00	250	195	1.058	
96.00	250	192	1.115	
120.00	250	182	1.308	
144.00	250	180	1.346	
168.00	250	176	1.423	
192.00	250	169	1.558	
216.00	250	171	1.519	
240.00	250	161	1.712	
0.00	500	500	0.000	
0.02	500	466	0.654	
24.00	500	431	1.327	
48.00	500	431	1.327	
72.00	500	430	1.346	
96.00	500	428	1.385	
120.00	500	421	1.519	
144.00	500	410	1.731	
168.00	500	410	1.731	
192.00	500	409	1.750	
216.00	500	398	1.962	
240.00	500	378	2.346	
0.00	1250	1250	0.000	
0.02	1250	1215	0.673	
18.50	1250	1181	1.327	
64.50	1250	1170	1.539	
117.00	1250	1117	2.558	
232.50	1250	1064	3.577	
554.00	1250	1013	4.558	

Table A.6 Chromium(VI) column experiment 1 — Experimental data.

APPENDIX C. Adsorption and Desorption of 50 mg/l Cr(VI) in Distilled Water pore volume 18.5 ml -----Cr(VI) removed from soil Sample Cr(VI) Cr (VI) Volume Pore vol. Cum. pv pH cum Cr(VI) point cum mq/1mole mL mL mL mol mol cmol/kg ug/g C 1 C 2 C 3 C 4 C 5 0.00 0.0E+00 25.0 1.35 1.35 3.85 0.0E+00 0.0E+00 2.22 3.57 3.84 9.7E-07 9.7E-07 0.000 0.00 1.23 2.4E-05 41.0 0.002 1.01 1.2E-05 14.50 2.8E-04 40.5 2.19 5.76 3.82 1.1E-05 0.025 11.75 31.10 6.0E-04 38.0 2.05 7.81 2.3E-05 3.72 3.5E-05 0.070 23.64 40.85 7.9E-04 1.5E-04 142.5 7.70 15.51 1.1E-04 3.72 0.294 116.42 46.44 8.9E-04 95.0 5.14 20.65 3.77 8.5E-05 2.3E-04 0.464 88.24 C 7 46.00 8.8E-04 126.0 6.81 27.46 3.83 1.1E-04 3.4E-04 0.687 115.92 CD8 1.84 8.9E-04 46.44 34.0 29.30 3.0E-05 3.86 3.7E-04 0.747 31.58 CD9 43.64 8.4E-04 14.0 0.76 30.05 3.76 1.2E-05 3.9E-04 0.771 12.22 CD10 \* 4.13 4.4E-04 39.45 7.6E-04 66.0 3.57 33.62 5.0E-05 0.871 52.07 CD11 \* 3.3E-04 17.11 23.0 1.24 34.86 4.40 7.6E-06 4.4E-04 0.886 7.87 CD12 \* 5.57 1.1E-04 138.5 7.49 42.35 4.18 1.5E-05 4.6E-04 0.916 15.43 CD13 \* 2.64 5.1E-05 101.0 5.46 47.81 5.1E-06 4.6E-04 4.55 0.926 5.33 CD14 \* 1.41 2.7E-05 165.5 8.95 56.76 4.5E-06 4.66 4.7E-04 0.935 4.67 CD15 \* 8.38 0.74 1.4E-05 155.0 65.14 4.83 2.2E-06 4.7E-04 0.939 2.29 CD16 \* 0.47 9.0E-06 171.5 9.27 74.41 4.66 1.6E-06 4.7E-04 0.942 1.61 CD17 \* 0.37 7.1E-06 174.0 9.41 83.81 5.16 1.2E-06 4.7E-04 0.945 1.29 CD18 \* 4.6E-06 0.24 199.0 10.76 94.57 5.25 9.2E-07 4.7E-04 0.947 0.96 CD19 \* 0.21 4.0E-06 133.0 7.19 101.76 5.4E-07 4.80 4.7E-04 0.948 0.56 CD20 \* 3.3E-06 9.73 111.49 0.17 180.0 5.9E-07 4.7E-04 5.16 0.949 0.61 CD21 \* 0.13 2.5E-06 204.0 11.03 122.51 5.18 5.1E-07 4.8E-04 0.950 0.53 CD22 \* 0.11 2.1E-06 237.0 12.81 135.32 5.0E-07 4.8E-04 0.951 5.15 0.52 CD23 \* 0.10 1.9E-06 141.0 7.62 142.95 5.20 2.7E-07 4.8E-04 0.952 0.28 CD24 \* 0.09 1.7E-06 150.0 8.11 151.05 5.23 2.6E-07 4.8E-04 0.952 0.27 CD25 \* 1.3E-06 2.7E-07 0.07 11.03 162.08 204.0 5.26 4.8E-04 0.953 0.29 CD26 \* 0.07 1.3E-06 213.0 11.51 173.59 5.14 2.9E-07 4.8E-04 0.953 0.30 CD27 \* 0.08 1.5E-06 115.0 6.22 179.81 5.54 1.8E-07 4.8E-04 0.954 0.18 CD28 \* 0.07 1.3E-06 110.0 5.95 185.76 5.52 1.5E-07 4.8E-04 0.954 0.15 \* Desorption with distilled water.

Table A.7 Chromium(VI) column experiment 2 — Experimental data.

APPENDIX B. Adsorption and Desorption of 10 mg/l Cr(VI) in Distilled Water pore volume 18.5 ml \_\_\_\_\_ Cr(VI) removed from soil Sample Cr(VI) Cr (VI) Volume Pore vol. Cum. pv pH Cr(VI) point cum mg/l mole mL mL mL mol mol cmol/kg ug/g \_\_\_\_\_ 0.00 0.0E+00 LAR 1 27.0 1.46 1.46 4.20 0.0E+00 0.0E+00 2.92 3.63 0.0E+00 0.0E+00 4.57 3.68 0.0E+00 0.0E+00 AB 2 0.00 0.0E+00 27.0 1.46 0.00 0.00 ав з 0.00 0.0E+00 30.5 1.65 0.00 0.00 AB 4 0.00 0.0E+00 20.0 1.08 5.65 3.71 0.0E+00 0.0E+00 0.00 0.00 AB 5 AB 6 0.00 0.0E+00 23.0 1.24 6.89 3.82 0.0E+00 0.0E+00 0.00 0.00 3.21 6.2E-05 95.0 5.14 12.03 5.9E-06 3.85 5.9E-06 0.01 6.10 AB 7 6.54 1.3E-04 64.0 1.4E-05 3.46 15.49 3.90 8.0E-06 0.03 8.37 AB 8 7.53 1.4E-04 37.0 2.00 17.49 3.85 5.4E-06 1.9E-05 0.04 5.57 AB 9 AB10 7.85 1.5E-04 31.0 1.68 19.16 3.88 4.7E-06 2.4E-05 0.05 4 87 7.88 1.5E-04 136.0 7.35 26.51 3.92 2.1E-05 4.5E-05 0.09 21.43 AB11 8.39 1.6E-04 88.5 4.78 31.30 3.68 1.4E-05 5.9E-05 0.12 14.85 AB12 AB13 8.53 1.6E-04 121.5 6.57 37.86 3.93 2.0E-05 7.9E-05 0.16 20.73 8.92 1.7E-04 27.0 1.46 39.32 3.93 4.6E-06 8.3E-05 0.17 4.82 ABD14\* 9.00 1.7E-04 5.0 0.27 39.59 3.93 8.7E-07 8.4E-05 0.17 0.90 ABD15\* 82.5 8.70 1.7E-04 14.35 4.46 44.05 4.30 1.4E-05 9.8E-05 0.20 ABD16\* 3.84 7.4E-05 143.5 7.76 51.81 4.49 1.1E-05 1.1E-04 0.22 11.02 ABD17\* 1.61 3.1E-05 111.0 6.00 57.81 4.51 3.4E-06 1.1E-04 0.22 3.57 ABD18\* 1.04 2.0E-05 156.5 8.46 66.27 4.30 3.1E-06 1.2E-04 0.23 3.26 1.6E-06 1.1E-06 ABD19\* 0.56 1.1E-05 150.0 8.11 74.38 4.49 1.2E-04 0.23 1.68 ABD20\* 0.37 7.1E-06 159.0 8.59 82.97 4.45 1.2E-04 0.24 1.18 ABD21\* 0.27 5.2E-06 181.5 9.81 92.78 4.47 9.4E-07 1.2E-04 0.24 0.98 ABD22\* 0.20 3.8E-06 186.0 10.05 102.84 7.2E-07 4.92 1.2E-04 0.24 0.74 ABD23\* 0.16 3.1E-06 148.0 8.00 110.84 4.74 4.6E-07 1.2E-04 0.24 0.47 ABD24\* 2.5E-06 0.13 168.0 9.08 119.92 5.09 4.2E-07 1.2E-04 0.24 0.44 ABD25\* ABD26\* 0.12 2.3E-06 193.0 10.43 130.35 4.96 4.5E-07 1.2E-04 0.24 0.46 0.11 2.1E-06 231.0 12.49 142.84 4.95 4.9E-07 1.2E-04 0.24 0.51 ABD27\* 1.9E-06 165.0 8.92 151.76 0.10 5.06 3.1E-07 1.2E-04 0.24 0.32 ABD28\* 0.06 1.2E-06 142.0 7.68 159.43 4.86 1.6E-07 1.2E-04 0.24 0.17 ABD29\* 0.06 1.2E-06 200.0 10.81 170.24 2.3E-07 5.05 0.24 1.2E-04 0.24 ABD30\* 0.05 9.6E-07 211.0 11.41 181.65 5.30 2.0E-07 1.2E-04 0.24 0.21 ABD31\* 0.06 1.2E-06 115.0 6.22 187.86 5.32 1.3E-07 1.2E-04 0.24 0.14 ABD32\* 0.05 9.6E-07 135.0 7.30 195.16 4.75 1.3E-07 1.2E-04 0.25 0.14 \* Desorption with Distilled water

Table A.8 Chromium(VI) column experiment 3 — Experimental data.

APPENDIX G. Adsorption of 10 mg/l Cr(VI) in Distilled Water and Desorption with 0.02 M

KH2PO4 at pH 4.53 and with pore volume of 18.4 ml.

Sample no	Cr(VI) mg/l	Cr (VI) mole	Volume mL	Pore vol.	Cum. pv mL	рН	Cr(VI) re point mol	emoved fro	om soil cum cmol/kg	Cr(VI)
D 1 D 2	0.00			1.43 3.57		3.89		0.0E+00		0.00
<b>D</b> 3	2.86	5.5E-05	157.0	8.49	13.49	3.96 3.96	8.6E-06	0.0E+00 8.6E-06	0.017	0.00 8.98
D 4 D 5	7.32 8.32		85.5 166.0	4.62 8.97	18.11 27.08	3.89 4.30	1.2E-05 2.7E-05	2.1E-05 4.7E-05	0.041 0.094	12.52 27.62
D 6 D 7	9.42 9.32		112.0 168.5	6.05 9.11	33.14 42.24	4.28	2.0E-05 3.0E-05	6.8E-05 9.8E-05	0.135 0.195	21.10 31.41
D 8	9.34	1.8E-04	17.0 56.5	0.92 3.05	43.16	4.30	3.1E-06 1.0E-05	1.0E-04 1.1E-04	0.202	3.18
DP10 *	9.55	1.8E-04	14.5	0.78	47.00	3.84	2.7E-06	1.1E-04	0.227	2.77
DP11 * DP12 *	10.84 8.06	2.1E-04 1.6E-04	37.0 154.5	2.00 8.35	49.00 57.35	4.14 3.94	7.7E-06 2.4E-05	1.2E-04 1.5E-04	0.243 0.291	8.02 24.91
DP13 * DP14 *	0.11 0.04	2.1E-06 7.7E-07	168.5 195.0	9.11 10.54	66.46 77.00	3.98 3.99	3.6E-07 1.5E-07	1.5E-04 1.5E-04	0.291 0.292	0.37 0.16
DP15 * DP16**	0.00 0.00		124.0 246.0	6.70 13.30	83.70 97.00	3.97 3.67	0.0E+00 0.0E+00	1.5E-04 1.5E-04	0.292	0.00

<sup>\*</sup> Desorption with 0.02 M KH2PO4 at pH 4.53. \*\* Desorption with 0.1 M KOH at pH 9.0.

Table A.9 Chromium(VI) one-hour isotherm — Experimental data.

•	Ci,			s,	C,	C/S,	
ppm:	umo1/1:	рн:	cmol/kg:	umo1/kg:			
2	38.46		0.0248				
5	96.16	3.00	0.0403	403.00	80.04	0.20	
10	192.32	3.00	0.0403 0.0624	624.00	167.36	0.27	
20	384.64	3.00	0.0778	778.00	353.52	0.45	
50	961.61	3.00	0.3000	3000.00	841.61	0.28	
Ci,	Ci, umol/l:		s,	s,	C,	C/S,	
ppm:	umol/1:	pH:	cmol/kg:	umol/kg:	umol/1:	Kg/L:	
2	38.46		0.0219				
5	96.16	4.00	0.0345	345.00	82.36	0.24	
10	192.32	4.00	0.0545	545.00	170.52	0.31	
20	384.64	4.00	0.0637 0.2160	637.00	359.16	0.56	
50	961.61	4.00	0.2160	2160.00	875.21	0.41	
Ci,	Ci,		s,	s,	C,	c/s,	
ppm:	umo1/1:	pH:	cmol/kg:	umol/kg:	umol/l:	Kg/L:	
2	38.46	5.00	0.0190	190.00	30.86	0.16	
5	96.16	5.00	0.0287	287.00	84.68	0.30	
10	192.32	5.00	0.0466	466.00	173.68	0.37	
	384.64	5.00	0.0496				
50	961.61	5.00	0.1160	1160.00	915.21	0.79	
Ci,	Ci,		s,	s,	C,	c/s,	
ppm:	umo1/1:	pH:	cmol/kg:	umol/kg:			
	38.46	6.00	0.0160	160.00	32.06	0.20	
5	96.16	6.00	0.0229	229.00	87.00	0.38	
10	192.32	6.00	0.0387 0.0356	387.00	176.84	0.46	
20	384.64	6.00	0.0356	356.00	370.40	1.04	
50	961.61	6.00	0.1000	1000.00	921.61	0.92	

# APPENDIX B

# FINITE DIFFERENCE PROGRAM

The following is a complete listing of the finite difference program developed to generate soil column breakthrough curves. The program consists of the following subroutines:

-	Main.Bas	(main subroutine)
-	Inp.Bas	(input subroutine)
-	Scont.Bas	(screen control subroutine)
-	Initial.Bas	(initializing subroutine)
-	Calc.Bas	(calculation subroutine)
-	Boundary.Bas	(boundary subroutine)
-	Estimate.Bas	(estimation subroutine)
-	Out.Bas	(output subroutine)

### Main.Bas Subroutine

```
Main.Bas
                 This is the Main program
It dimensions the 1-D arrays used in the program.
                 It calls up the subroutines.
                 It also links the subroutines at compile time.
                 IMPORTANT:
                 Before compiling the program, this subroutine has to be declared "Main File" in the "File" options of the TURBO BASIC environment !!!
 ***** 1-D arrays are dimensioned *****
    DIM Valence#(10), b#(10), Cinitial#(10), Cdesorb#(10), Csoil#(10), MW#(10)
    DIM LastnewMOB#(10), LastnewIMMOB#(10), SoldMOB#(10), SoldIMMOB#(10), SnewMOB#(10), SnewIMMOB#(10), PrecalcMOB#(10), PrecalcIMMOB#(10),_
         EfunctionMOB#(10), EfunctionIMMOB#(10), SlopeMOB#(10), SlopeIMMOB#(10),
        Effluent#(10)
***** Subroutines are linked together *****
$INCLUDE "Inp.Bas"
$INCLUDE "Scont.Bas"
$INCLUDE "Initial.Bas"
$INCLUDE "Calc.Bas"
$INCLUDE "Boundary.Bas"
$INCLUDE "Estimate.Bas"
$INCLUDE "Out.Bas"
****** Subroutines are called serially during program performance *****
CALL Inputsub
CALL Screencontrol
CALL Initialize
CALL Calculation
CALL Outputfile
CLS
PRINT
PRINT "Program terminated"
END
```

### Inp.Bas Subroutine

```
INP.BAS
             This is the Input subroutine. It allows to create a new input
             file as well as to use and/or modify an already existing file.
             Finally one can also save the created input file.
             The sub also dimensions the 2-D arrays
SUB Inputsub
   ***** Variable declarations *****
  SHARED SpeciesNo%, I%, CellNo%, Stepchoice%
  SHARED Dispersivity#, Diameter#, Density#, WatercontMOB#, WatercontIMMOB#, _ Alpha#, DX#, Totaltime#, Timestep#, Tdesorb#, Convergence#, F#, Q#__
         Flowrate#
  SHARED Valence#(), b#(), Cinitial#(), Cdesorb#(), Csoil#(), MW#(), ColdMOB#()_
         ColdIMMOB#(), CnewMOB#(), CnewIMMOB#(), Cellarray#()
  SHARED Outputfilename$
  LOCAL Previous%, Ok%, Savefile%, Locpos%, Sstop%, Choice%, SStep%
  LOCAL Oldinputfilename$, Newinputfilename$
  CLS
  PRINT "Would you like to use and/or modify an already existing input file ?"
  PRINT
  PRINT "Yes --> hit <1> <enter> , No --> hit <0> <enter> ";
  INPUT Previous%
   '***** the following edits/modifies an already existing file *****
  IF Previous% = 1 THEN
     PRINT
     PRINT "Specify the drive and the filename."
     PRINT "Make sure the disk is in your specified drive and ready to go !"
     PRINT
     INPUT "Input filename ";Oldinputfilename$
      '**** here the data file is read into memory *****
     OPEN Oldinputfilename$ FOR INPUT AS #1
        INPUT #1, SpeciesNo%
        FOR I% = 1 TO SpeciesNo%
           INPUT #1, MW#(I%), Valence#(I%), b#(I%), Cinitial#(I%), Cdesorb#(I%), Csoil(I%)
        NEXT 1%
        INPUT #1, Dispersivity#, Diameter#, Density#, Q#, F#, WatercontMOB#,__
WatercontIMMOB#, Alpha#, CellNo%, Dx#, Totaltime#, Timestep#_
                Tdesorb#, Convergence#, Stepchoice%, Flowrate#
     CLOSE #1
     0k\% = 0
     Choice% = 0
     DO
        ***** here the data file is displayed on screen *****
             CLS
        LOCATE 1.15
        PRINT "DATASHEET of ";Oldinputfilename$
        PRINT
        PRINT "Spec-No. MW
                                    Valence
                                               Langmuir coeff. Cinput(L) Cdesorb(L) Cinit(SL)"
        PRINT "
                         (g/mol)
                                             (L/mmol)
                                                              (mg/L)
                                                                          (mg/L)
                                                                                       (mg/L)
        PRINT
        FOR I% = 1 TO SpeciesNo%
           Locpos\% = 5 + I\%
           LOCATE Locpos%,3
           PRINT USING "##";1%
           LOCATE Locpos%,10
           PRINT USING "###";MW#(I%)
           LOCATE Locpos%,21
           PRINT USING "##"; Valence#(I%)
           LOCATE Locpos%,31
```

```
PRINT USING "#####.##";b#(1%)
    LOCATE Locpos%,48
    PRINT USING "####.##"; Cinitial#(1%)
    LOCATE Locpos%,60
    PRINT USING "####.##";Cdesorb#(I%)
    LOCATE Locpos%,72
    PRINT USING "####.##"; Csoil#(1%)
 NEXT 1%
 LOCATE 11.1
 PRINT "Aquifer / Soil parameters:
                                                      General parameters:"
 PRINT
 PRINT "Dispersivity
PRINT USING "###.##";Dispersivity#
PRINT "Bulk density (g/cm^3)";
 PRINT USING "###.##"; Density#
PRINT "Max. adsorp. cap. (umol/g)";
PRINT USING "###.##";Q#
PRINT "Soil-Frac.(f) -> mobile H2O";
PRINT USING "###.##";F#
PRINT "Water content (mob)
 PRINT USING "###.##"; WatercontMOB#
PRINT "Water content (immob) ";
PRINT USING "####"; WatercontIMMOB#
PRINT "Transfer coeff. Alpha (1/h)";
PRINT USING "###.##";Alpha#
PRINT "Flowrate (cm^
                                (cm^3/n)";
PRINT USING "###.##";Flowrate#
PRINT "Column diameter
                                     (cm)";
PRINT USING "###.##"; Diameter#
 LOCATE 15,40
PRINT "Number of cells
                                          ":CellNo%
 LOCATE 16.40
PRINT "Length of one cell
PRINT USING "###.####";Dx#
                                    (cm)";
LOCATE 17,40
PRINT "Total calculation time (h)";
PRINT USING "###.####";Totaltime#
LOCATE 18,40
PRINT "Length of time step (h)";
PRINT USING "###.###";Timestep#
LOCATE 19,40
PRINT "Desorption start time (h)";
PRINT USING "###.###";Tdesorb#
LOCATE 20,40
PRINT "Convergence criteria (mmol)";
PRINT USING "^^^^";Convergence#
LOCATE 21,40
PRINT "Output interval:
                                   each ";Stepchoice%;"steps"
LOCATE 25,1
INPUT "To continue, hit <enter> ";Sstop%
       ***** display of data change option menu *****
PRINT "
                                        ******
PRINT "
                                          OPTION MENU
PRINT "
PRINT
PRINT "
                                   Change the following parameters:"
PRINT
PRINT " 1. Number of species (max. 10)
                                                                   13. Number of cells"
PRINT " 2. Valence / MW of a species
                                                                   14. Length of one cell"
PRINT " 3. Langmuir coeff. of a species
PRINT " 4. Initial concentrations of a species
PRINT "
                                                                 15. Total calculation time"
                                                               16. Length of time step"
                                                                17. Desorption time"
PRINT " 5. Dispersivity
PRINT " 6. Column diameter
PRINT " 7. Bulk density
                                                                18. Convergence criteria"
                                                                  19. Output Interval"
                                                                 20. Flowrate"
PRINT " 8. Adsorption capacity"
PRINT " 9. Fraction in contact w mobile H2O"
PRINT "10. Water content (mobile)"
PRINT "11. Water content (immobile)"
```

```
PRINT "12. Transfer coefficient Alpha"
PRINT
PRINT "Choose by hitting <number> <enter>. To exit Option Menu, hit <enter>"
INPUT "Your choice "; Choice%
     '***** select case carries out the chosen change option *****
SELECT CASE Choice%
  CASE 0
     0k\% = 1
  CASE 1
     PRINT "Number of species:";SpeciesNo%
     INPUT "New value (max. 10) ";SpeciesNo%
     PRINT
  CASE 2
     INPUT "For which species do you want to change the valence / mol. weight ";1% PRINT "Species No: ";1%;" Valence: ";Valence#(1%);" Mol. weight: ";MW#(1%) INPUT "New Valence ";Valence#(1%)
     INPUT "New Molecular Weight ":MW#(1%)
  CASE 3
     INPUT "For which species do you want to change the Langmuir coeff."; I%
     PRINT "Species No: ";1%;" Langmuir coeff.: ";b#(1%)
     INPUT "New Langmuir coeff. ";b#(I%)
  CASE 4
     INPUT "For which species do you want to change init. concentrations ";1%
     PRINT "Species No: ";1%
     PRINT "Initial liquid concentr.
PRINT "Desorption concentr.
                                     ";Cinitial#(1%)
                                        ";Cdesorb#(1%)
     PRINT "Init. soil-Liquid conc.
                                      ":Csoil#(1%)
     PRINT
     INPUT "New init. liquid conc. ";Cinitial#(1%)
     INPUT "New desorption conc. ";Cdesorb#(1%)
     INPUT "New init. soil-liquid conc. ";Csoil#(1%)
  CASE 5
     PRINT "Dispersivity
                              : ";Dispersivity#
     INPUT "New value ";Dispersivity#
  CASE 6
     PRINT "Column diameter : ";Diameter#
     INPUT "New value
                          ";Diameter#
  CASE 7
    PRINT "Bulk density : ";Density# INPUT "New value ";Density#
  CASE 8
     PRINT "Max. adsorption capacity: ";Q#
     INPUT "New value"; O#
  CASE 9
    PRINT "Soil fraction (f) in contact with mobile water: ":F#
    INPUT "New value ";F#
  CASE 10
     PRINT "Water content (mobile): "; WatercontMOB#
    INPUT "New value ": WatercontMOB#
  CASE 11
    PRINT "Water content (immobile): "; WatercontIMMOB#
    INPUT "New value "; WatercontIMMOB#
  CASE 12
    PRINT
            "Transfer coefficient Alpha: "; Alpha#
    INPUT "New value ";Alpha#
  CASE 13
    PRINT '
             'Number of cells: ";CellNo%
    INPUT " New value ";CellNo%
 CASE 14
    PRINT "Length of one cell : ";Dx#
    INPUT "New value ";Dx#
 CASE 15
    PRINT "Total calculation time: ":Totaltime#
    INPUT "New value "; Totaltime#
 CASE 16
    PRINT "Length of time step: ";Timestep#
    INPUT "New value ";Timestep#
 CASE 17
    PRINT "Desorption time : ";Tdesorb#
```

```
INPUT "New value ";Tdesorb#
      CASE 18
         PRINT "Convergence criteria: ";Convergence#
         INPUT "New value ";Convergence#
      CASE 19
         PRINT "Output intervall: ";Stepchoice%;" timesteps"
         INPUT "New value ";Stepchoice%
      CASE 20
         PRINT "Flowrate (cm^3/h)
                                         :":Flowrate#
         INPUT "New value
                                ":Flowrate#
   END SELECT
LOOP UNTIL Ok\% = 1
PRINT
CLS
INPUT "Specify drive and name of OUTPUT data file: ";Outputfilename$
***** here a whole new data file can be entered *****
0k\% = 0
DO
   CLS
   INPUT "Number of species to include into model (max. 10) ";SpeciesNo%
   PRINT
   PRINT
   CLS
   FOR 1% = 1 TO SpeciesNo%
      PRINT "Species No. :
                                                                ":I%
      INPUT "Valence
INPUT "Molecular Weight
                                                                ";Valence#(I%)
                                                         (g/mol) ":MW#(I%)
      INPUT "Langmuir coefficient b ((L/mmol) Valence)";b#(I%)
INPUT "Input liquid concentration (mg/L)";Cinitial#(I%)
      INPUT "Input liquid concentration INPUT "Desorption - Input liquid conc.
                                                      (mg/L) ";Cdesorb#(I%)
(mg/L) ";Csoil#(I%)
      INPUT "Initial liquid conc. in soil column
      PRINT
   NEXT 1%
   PRINT
   PRINT "Is the input correct? If yes, hit <1> <enter>"; INPUT Ok%
LOOP UNTIL Ok\% = 1
0k\% = 0
DO
   CLS
   PRINT "Enter the following AQUIFER/SOIL parameters:"
   PRINT
   INPUT "Dispersivity
                                                                      ";Dispersivity#
                                               (cm)
                                                (g/cm^3)
   INPUT "Bulk density
                                                                       ";Density#
   INPUT "Max. adsorption capacity
INPUT "Fraction of solids in contact w/
                                                 (umol/g)
                                                                        ";Q#
                                               mobile water (0-1)
                                                                      ":F#
                                                                         ";WatercontMOB#
";WatercontIMMOB#
   INPUT "Water content (mobile)
                                                 (cm^3/cm^3)
   INPUT "Water content (immobile)
INPUT "Transfer coeff. ALPHA
                                                 (cm^3/cm^3)
                                                                        ";Alpha#
                                                  (1/h)
   INPUT "Flowrate
                                                (cm^3/h)
                                                                       ":Flowrate#
   PRINT
   PRINT "Is the input correct? If yes, hit <1> <enter>";
   INPUT Ok%
LOOP UNTIL Ok% = 1
0k\% = 0
DO
   CLS
   PRINT "Enter the following general parameters:"
   PRINT
   INPUT "Number of cells
INPUT "Length of one cell
INPUT "Column diameter
                                                               ":CellNo%
                                                     (cm)
                                                                ";Dx#
                                                                 ";Diameter#
                                                       (cm)
   INPUT "Total Calculation Time
                                                                ':Totaltime#
                                                      (h)
   INPUT "Length of time step
INPUT "Time when DESORPTION begins
INPUT "Convergence criteria
                                                               ";Timestep#
                                                      (h)
                                                                      ":Tdesorb#
                                                             (ከ)
                                                                ";Convergence#
                                                  (mmol)
                                               (timesteps)
   INPUT "Frequency of output
                                                               ";Stepchoice%
   INPUT "Specify drive and name for OUTPUT data file: ";Outputfilename$
```

```
PRINT
        PRINT "Is the input correct? If yes, hit <1> <enter>";
        INPUT Ok%
     LOOP UNTIL 0k\% = 1
   ***** end of data input or modification *****
   END IF
   ***** 2-D arrays are dimensioned here *****
   DIM ColdMOB#(SpeciesNo%+1,CellNo%+1), ColdIMMOB#(SpeciesNo%,CellNo%),_
   CnewMOB#(SpeciesNo%,CellNo%), CnewIMMOB#(SpeciesNo%,CellNo%)
SStep% = INT(Totaltime# / (Timestep# * Stepchoice%)) + 1
   DIM Cellarray#(SStep%, SpeciesNo%+1)
   CLS
   ***** the following contains the option to save the current data file *****
   Savefile% = 0
   PRINT "Now you have entered all necessary input data !"
   PRINT
   PRINT "You have now the option to save all this data in a file ..."
INPUT "If you wish to do this, hit <1> <enter>";Savefile%
   IF Savefile% = 1 Then
     CLS
     PRINT
     PRINT "Please prepare your disk and make sure everything is ready to go !"
     PRINT
     PRINT "Please specify the drive and the filename.
     PRINT
     PRINT "* DANGER !!! If a file with the same name already exists, the new *"
     PRINT
     INPUT "Input Filename
                                  ";Newinputfilename$
     '**** here the data is saved in a file ***
     ***** the format has to be consistent with the data read section above *****
     OPEN Newinputfilename$ FOR OUTPUT AS #1
        WRITE #1, SpeciesNo%
        FOR I% = 1 TO SpeciesNo%
           WRITE #1, MW#(I%), Valence#(I%), b#(I%), Cinitial#(I%), Cdesorb#(I%), Csoil#(I%)
        NEXT 1%
        WRITE #1, Dispersivity#, Diameter#, Density#, Q#, F#, WatercontMOB#_
WatercontIMMOB#, Alpha#, CellNo%, Dx#, Totaltime#, Timestep#,_
                Tdesorb#, Convergence#, Stepchoice%, Flowrate#
     CLOSE #1
   END IF
END SUB
```

# Scont.Bas Suroutine

```
' Scont.Bas
' This is the Sreen control sub. It allows the user to tell, which part of the program is being processed currently.
' The sub gets its input from "locate" and "print" statements within each sub.

SUB Screencontrol

CLS
PRINT PRINT PRINT
PRINT "Program is currently processing: "
PRINT
PRINT PRINT
PRINT
PRINT
PRINT
PRINT "Current stepnumber:"
PRINT
END SUB
```

#### Initial.Bas Subroutine

```
Initial.Bas
             This sub initializes the cells (using Csoil(1%))
             It also performs the precalculation of constant parameters
             used in the calculation sub.
Sub Initialize
       '***** variable declarations *****
             SHARED SpeciesNo%, I%, CellNo%, N%
             SHARED ColdMOB#(), ColdIMMOB#(), Cinitial#(), Cdesorb#(), Csoil#(), MW#()_
             SHARED Dispersivity#, Porevelocity#, Dx#, Timestep#, F#, Alpha#, Density#,...
             WatercontMOB#, WatercontIMMOB#, Value1#, Value2#, Value3#, Value4#,
              Value5#, Value6#, Totalvolume#, Flowrate#, Diameter#
             LOCAL Disp#, Dispersion#, Pi#
   ***** statement for screen control sub *****
  LOCATE 3,36
  PRINT "Initialize Subroutine"
   ***** mass conc. are converted to molar conc.; if the initial conc. of a species
         is zero, it gets the conc.1E-100 assigned to allow for division later to
         convert to relative concentrations *****
  FOR I% = 1 TO SpeciesNo%
     Cinitial#(I\%) = Cinitial#(I\%) / MW#(I\%)
     Cdesorb#(I\%) = Cdesorb#(I\%) / MW#(I\%)
     Csoil#(I\%) = Csoil#(I\%) / MW#(I\%)
     IF Csoil#(1%) = 0 THEN Csoil#(1%) = 1E-100
  NEXT 1%
  '***** cell conc. are initialized *****
  FOR N% = 1 TO CellNo%
     FOR I% = 1 TO SpeciesNo%
        ColdMOB\#(I\%,N\%) = Csoil\#(I\%)
        ColdIMMOB\#(1\%,N\%) = Csoil\#(1\%)
     NEXT 1%
  NEXT N%
   ***** calculation of average pore velocity, column vol. and dispersion coeff.;
         dispersion coeff. is corrected for numerical disp. error and
         is not allowed to be smaller than zero *****
  Pi\# = 4 * ATN(1)
  Porevelocity# = Flowrate# / ((Diameter#/2)^2 * Pi# * WatercontMOB#)
  Totalvolume# = (Diameter#/2)^2 * Pi# * CellNo% * Dx#
  Dispersion# = Porevelocity# * Dispersivity#
  Disp# = Dispersion# - Porevelocity# * (Dx# - Porevelocity# * Timestep#) / 2
  IF \hat{D}isp\# < \hat{0} THEN Disp\# = 0
  ***** constant fractions for FD equations are computed and
        displayed on screen ***
  Value 1# = Disp# * Timestep# / (Dx# * Dx#)
  Value2# = Porevelocity# * Timestep# / Dx#
  Value3# = F# * Density# / WatercontMOB#
  Value4# = Alpha# * Timestep# / WatercontMOB#
  Value5# = (1 - F#) * Density# / WatercontIMMOB#
  Value6# = Alpha# * Timestep# / WatercontIMMOB#
 LOCATE 10,1
 PRINT "Dispersion factor
 PRINT USING "########"; Value1#
PRINT "Advection factor : ";
 PRINT USING "#####.####":Value2#
 PRINT "Sorption factor (mob): ";
 PRINT USING "#####.#####";Value3#
PRINT "Transfer factor (mob): ";
 PRINT USING "#####.####"; Value4#
 PRINT "Sorption factor (im):
 PRINT USING "#####.#####";Value5#
```

```
PRINT "Transfer factor (im): ";
PRINT USING "########"; Value6#
'***** absolute of valence in case of negative input value *****
FOR I% = 1 TO SpeciesNo%
Valence#(I%) = ABS(Valence#(I%))
NEXT I%
END SUB
```

### Calc.Bas Subroutine

```
Calc.Bas
            This sub performs the finite difference calculations.
            It contains a convergence loop for each cell within a time
            The results of the calculations are written to a 2-D array
SUB Calculation
   '***** variable declarations *****
  SHARED I%, SpeciesNo%, N%, CellNo%, Flag1%
  SHARED ColdMOB#(), ColdIMMOB#(), CnewMOB#(), CnewIMMOB#(),
        Cellarray#(), LastnewMOB#(), LastnewIMMOB#(), b#(), Valence#(),
         SoldMOB#(), SoldIMMOB#(), SnewMOB#(), SnewIMMOB#(), Cinitial#(),_
        Cdesorb#(), PrecalcMOB#(), PrecalcIMMOB#(), EfunctionMOB#(), EfunctionIMMOB#(), SlopeMOB#(), SlopeIMMOB#(), Effluent#()
  SHARED T#, Timestep#, Totaltime#, Q#, Convergence#, Value1#,
         Value2#, Value3#, Value4#, Value5#, Value6#, Stepchoice%, Stepcheck#
  LOCAL StepNo&
  LOCAL SumoldMOB#, SumoldIMMOB#, SumnewMOB#, SumnewIMMOB#,_
        Check#, Check1#, Check2#
   '**** statement for screen control sub *****
  LOCATE 3.36
  PRINT "Calculation Subroutine"
   ***** begin time loop *****
  StepNo&=0
  FOR T# = Timestep# TO Totaltime# STEP Timestep#
     INCR StepNo&, 1
     ***** statement for screen control *****
     LOCATE 6,22
     PRINT StepNo&
     ***** calling up boundary conditions *****
     Call Boundaries
     '**** begin cell loop *****
     FOR N% = 1 TO CellNo%
        ***** computing the sum of b*C(I) for old time *****
        SumoldMOB# = 0
        SumoldIMMOB# = 0
        FOR I% = 1 TO SpeciesNo%
          SumoldMOB# = SumoldMOB# + b#(I%) * ColdMOB#(I%,N%)^{}
          (1/Valence#(1%))
          SumoldIMMOB# = SumoldIMMOB# + b#(I%) * ColdIMMOB#(I%,N%)^_
          (1/Valence#(1%))
        NEXT 1%
        ***** computing solid phase conc. for all species at the old time using
              competitive Langmuir equation
              computing the explicit part of FD equations and saving the results
              in Precalc arrays ****
              in select case the upper boundary is applied (for cell 1) *****
        FOR I% = 1 TO SpeciesNo%
          SoldMOB#(I%) = (Q# * b#(I%) * ColdMOB#(I%,N%)^(1/Valence#(I%))) /_
          (1 + SumoldMOB#)
          SoldIMMOB#(I%) = (Q# * b#(I%) * ColdIMMOB#(I%,N%)^(1/Valence#(I%))) /_
          (1 + SumoldIMMOB#)
          SELECT CASE N%
             CASE >1
                PrecalcMOB\#(I\%) = ColdMOB\#(I\%,N\%) + Value1\# * (ColdMOB\#(I\%,N\%-1)_{-})
                - 2 * ColdMOB#(1%,N%) + ColdMOB#(1%,N%+1)) + Value2# *
                (ColdMOB#(1%,N%-1)-ColdMOB#(1%,N%)) - Value4# * (ColdMOB#(1%,N%)_
                 ColdIMMOB#(1%,N%))
             CASE 1
                PrecalcMOB#(I%) = ColdMOB#(I%,N%) + Value1# * (- ColdMOB#(I%,N%)_
                + ColdMOB#(I%,N%+1)) + Value2# * (ColdMOB#(I%,N%-1)
                 ColdMOB#(I%,N%)) - Value4#*(ColdMOB#(I%,N%) - ColdIMMOB#(I%,N%))
          END SELECT
          PrecalcIMMOB#(I%) = ColdIMMOB#(I%,N%) + Value6# * (ColdMOB#(I%,N%) -_
```

```
ColdIMMOB#(I%.N%))
 NEXT 1%
 Flag 1\% = 0
  ***** begin iteration loop within each cell *****
    ***** calling estimation sub, allowing for relaxation *****
    CALL Estimates
     ***** computing sum of b*C(I) at new time *****
    SumnewMOB# = 0
    SumnewIMMOB# = 0
    FOR I% = 1 TO SpeciesNo%
       SumnewMOB# = SumnewMOB# + b#(I%) * CnewMOB#(I%,N%)^_
       (1/Valence#(I%))
       SumnewIMMOB# = SumnewIMMOB# + b#(I%) * CnewIMMOB#(I%,N%)^_
       (1/Valence#(I%))
    NEXT 1%
     '**** computing solid phase conc. at new time using competitive
          Langmuir equation ****
    FOR 1% = 1 TO SpeciesNo%
       SnewMOB#(I%) = (Q# * b#(I%) * CnewMOB#(I%,N%)^(1/Valence#(I%))) /_
       (1 + SumnewMOB#)
       SnewIMMOB#(I%) = (Q# * b#(I%) * CnewIMMOB#(I%,N%)^(1/Valence#(I%))) /_
       (1 + SumnewIMMOB#)
    NEXT 1%
     ****** FD equations rewritten as zero functions (Efunction has to become zero)
          for both water phases; slope defines the derivative of the Efunctions
          with respect to the new conc. of a species (Cnew(I%,N%). Cnew
          of other species are treated as constants. *****
    FOR I% = 1 TO SpeciesNo%
       EfunctionMOB#(I%) = PrecalcMOB#(I%) - Value3# * (SnewMOB#(I%) -_
       SoldMOB#(I%)) - CnewMOB#(I%,N%)
       EfunctionIMMOB#(1%) = PrecalcIMMOB#(1%) - Value5# * (SnewIMMOB#(1%)_
       - SoldIMMOB#(I%)) - CnewIMMOB#(I%,N%)
SlopeMOB#(I%) = - Value3#*((1+SumnewMOB#)*Q#*b#(I%)*CnewMOB#(I%,N%)^_
       (1/Valence#(1%)-1)/Valence#(1%) - Q#*b#(1%)^2*CnewMOB#(1%,N%)^_
       (2/Valence#(I%)-1)/Valence#(I%))/(1+SumnewMOB#)^2 -1
       SlopeIMMOB#(I%) = - Value5#*((1+SumnewIMMOB#)*Q#*b#(I%)*CnewIMMOB#(I%.N%)^
       (1/Valence#(1%)-1)/Valence#(1%) - Q#+b#(1%)^2+CnewIMMOB#(1%,N%)^_
       (2/Valence#(I%)-1)/Valence#(I%))/(1+SumnewIMMOB#)^2 -1
     NEXT 1%
     ***** Newton-Raphson approximation *****
     FOR I% = 1 TO SpeciesNo%
       CnewMOB#(I%,N%) = LastnewMOB#(I%) - EfunctionMOB#(I%) / SlopeMOB#(I%)
       CnewIMMOB#(1%,N%) = LastnewIMMOB#(1%) - EfunctionIMMOB#(1%) / SlopeIMMOB#(1%)
     ***** checking the convergence criterium for all species *****
     Check# = 0
     FOR I% = 1 TO SpeciesNo%
       Check1# = ABS(EfunctionMOB#(I%))
       Check2# = ABS(EfunctionIMMOB#(1%))
       IF Check1# > Check# THEN
          Check# = Check1#
       ELSEIF Check2# > Check# THEN
          Check# = Check2#
       END IF
    NEXT 1%
  LOOP UNTIL Check# <= Convergence#
   ***** end of iteration loop within a cell *****
NEXT N%
'**** end of cell loop *****
***** check for very small conc.; to prevent values from running out of
     range they are set to zero; new conc. are assigned to old conc. *****
FOR N% = 1 TO CellNo%
  FOR I% = 1 TO SpeciesNo%
     IF CnewMOB#(1\%,N%) < 1*10^-300 THEN CnewMOB#(1\%,N%) = 0
     IF CnewIMMOB#(1%,N%) < 1*10^-300 THEN CnewIMMOB#(1%,N%) = 0
     ColdMOB\#(I\%,N\%) = CnewMOB\#(I\%,N\%)
     ColdIMMOB\#(I\%,N\%) = CnewIMMOB\#(I\%,N\%)
  Next I%
NEXT N%
```

```
'***** check for sampling time, if yes, conc. and time is written
' to Cellarray *****

Stepcheck# = StepNo& / Stepchoice%

IF Stepcheck# = INT(Stepcheck#) THEN

FOR I% = 1 TO SpeciesNo%

Cellarray#(Stepcheck#,I%) = Effluent#(I%)

NEXT I%

Cellarray#(Stepcheck#,SpeciesNo%+1) = T#

END IF

NEXT T#

'***** end of time loop *****

END SUB
```

# Boundary.Bas Subroutine

**END SUB** 

```
Boundary.Bas
              This sub sets the boundary conditions before a next time step
              is performed. The "FOR..NEXT" loop sets the lower boundary,
              whereas the "IF" loop sets the upper boundary initial conc.. This upper boundary is controlled by the time counter "T" and by the
              event "Tdesorb".
SUB Boundaries
        '**** variable declarations *****
              SHARED I%, SpeciesNo%, CellNo%
              SHARED ColdMOB#(), Cinitial#(), Cdesorb#(), Effluent#()
        SHARED T#, Tdesorb#
        LOCAL First%
        ***** set lower boundary for all species *****
        FOR I% = 1 TO SpeciesNo%
ColdMOB#(I%,CellNo%+1) = ColdMOB#(I%,CellNo%)
           Effluent#(1%) = ColdMOB#(1%,CellNo%+1)
        ***** set initial conc. for upper boundary
               the actual boundary restriction is in the Calculation sub *****
        First\% = 0
        IF T# < Tdesorb# THEN
FOR I% = 1 TO SpeciesNo%
              ColdMOB#(I%,First%) = Cinitial#(I%)
           NEXT 1%
        ELSE
           FOR I% = 1 TO SpeciesNo%
              ColdMOB#(I%,First%) = Cdesorb#(I%)
           NEXT I%
        END IF
```

### Estimate.Bas Subroutine

**END SUB** 

```
Estimate.Bas
              This sub calculates estimates for the new cell concentrations
              of all species in both the mobile and immobile layer. It also
              saves the previous (T-DT) concentrations in an array.
"CASE 0" is used for the initial (first) estimate of the cell
              concentrations."CASE 1" is used for estimates during the
              convergence procedure. This sub is being called by the "CALCULATION" sub.
              STATUS: use old conc. as estimates
SUB Estimates
       ***** variable declarations *****
              SHARED I%, SpeciesNo%, Flag1%, N%
SHARED ColdMOB#(), ColdIMMOB#(), CnewMOB#(), CnewIMMOB#()_
              LastnewMOB#(), LastnewIMMOB#()
             SELECT CASE Flag 1%
          CASE 0
             FOR I% = 1 TO SpeciesNo%
                CnewMOB\#(I\%,N\%) = ColdMOB\#(I\%,N\%)
                CnewIMMOB#(1%,N%) = ColdIMMOB#(1%,N%)
LastnewMOB#(1%) = ColdMOB#(1%,N%)
                LastnewIMMOB#(1%) = ColdIMMOB#(1%,N%)
                   NEXT 1%
                CASE 1
             FOR I% = 1 TO SpeciesNo%
                LastnewMOB#(I%) = CnewMOB#(I%,N%)
                LastnewIMMOB#(1%) = CnewIMMOB#(1%,N%)
             NEXT 1%
             END SELECT
       Flag 1% = 1
```

### Out.Bas Subroutine

END SUB

```
Out.Bas
            First the concentrations are converted to relative conc..
            Then the sub creates an output file (ASCII), which has
            been named previously (in the Input sub)
SUB Outputfile
       '**** variable declarations *****
            SHARED Cellarray#(), MW#(), Cinitial#()
            SHARED 1%, SpeciesNo%
            SHARED Outputfilename$
       SHARED Stepcheck#, Flowrate#, WatercontMOB#, Totalvolume#
            LOCAL Number%
  ***** statement for Screen control sub *****
  LOCATE 3,36
              Output Subroutine"
  PRINT "
  ***** set cinitial to allow for devision *****
  FOR I% = 1 TO SpeciesNo%
     IF Cinitial#(1%) = 0 THEN Cinitial#(1%) = 1E-100
  NEXT 1%
  Stepcheck# = INT(Stepcheck#)
  FOR Number% = 1 TO Stepcheck#
     ***** compute relative concentrations
           convert real time to mobile pore volumes *****
     FOR I% = 1 TO SpeciesNo%
        Cellarray#(Number%, 1%) = Cellarray#(Number%, 1%) / Cinitial#(1%)
     NEXT 1%
     Cellarray#(Number%,SpeciesNo%+1) = Cellarray#(Number%,SpeciesNo%+1) *_
     Flowrate# / (Totalvolume# * WatercontMOB#)
  NEXT Number%

'***** write to output data file *****
   OPEN Outputfilename$ FOR APPEND AS #1
   FOR Number% = 1 TO Stepcheck#
     PRINT #1, USING "####.###"; Cellarray#(Number%,SpeciesNo%+1);
     FOR I% = 1 TO SpeciesNo%
        PRINT #1, USING "###.###"; Cellarray#(Number%,I%);
     NEXT 1%
     PRINT #1, ""
  NEXT Number%
  CLOSE #1
```