Effects of vegetation canopy density and bank angle on near-bank patterns of turbulence 1 2 and Reynolds stresses 3 Nicole M. Czarnomski¹, Desireé D. Tullos², Robert E. Thomas³, and Andrew Simon⁴ 4 5 ¹Watershed Sciences Department, Utah State University, 5210 Old Main Hill, Logan, UT 84322; 6 PH (435) 797-2459; FAX (435) 797-1871; email: nicole.czarnomski@usu.edu 7 8 9 ²Department of Biological and Ecological Engineering, Oregon State University, Corvallis, OR 10 97331; PH (541) 737-2038; FAX (541) 737-2082; email: desiree.tullos@oregonstate.edu 11 12 ³Department of Geography, University of Hull, Cottingham Road, Hull, HU6 7RX. UK; PH +44 13 (0)1482 465355; FAX +44 (0)1482 466340; email: r.e.thomas02@members.leeds.ac.uk 14 15 ⁴CardnoENTRIX, 1223 Jackson Avenue East, Suite 301, Oxford, MS 38655; PH (662) 236-16 6983; FAX (662) 281-9942; email: andrew.simon@cardno.com 17 18 19 Abstract 20 Vegetation growing on the surface of a streambank has been shown to alter the shear stresses 21 applied to the boundary, but basic questions remain regarding the influence of vegetation and 22 streambank configurations on near-bank hydraulics. In the present study, Froude-scaled flume 23 experiments were used to investigate how changes in vegetation density (ratio of frontal area to 24 channel area, including both stems and leaves) and bank surface angle influence near-bank 25 turbulence intensities ($RMS_{u,v,w}$) and Reynolds stresses (τ_{uv} and τ_{uw}) estimated using velocities

obtained with an acoustic Doppler velocimeter positioned beneath the canopy. Results illustrate
 how, with increasing vegetation density, turbulence intensities and Reynolds stresses decreased

along the sloped bank surface but increased at the base of the slope and within the main channel.

The steeper bank angle resulted in greater vertical stresses on the bank surface than the shallower angle, but lateral momentum exchange was larger than vertical exchange along the base of the slope, regardless of bank angle. Leaves were an important influence on near-bank turbulence intensities and Reynolds stresses, while the influence of bank slope was small relative to the influence of vegetation density.

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36 Introduction

37 Vegetation on the base of streambanks may deflect flow and reduce near-bank velocities 38 and shear stresses, but may also induce turbulence, elevate shear stresses and promote localized 39 scour along the base of the bank surface (Wilkerson 2007; Yang et al. 2007; Gorrick 2009; 40 Hopkinson and Wynn 2009). Vegetation also generates turbulence in the vertical plane at the 41 interface between the canopy and the free-stream (Yang et al. 2007; White and Nepf 2008; 42 Hopkinson and Wynn 2009; Zong and Nepf 2010). Shear layers form at interfaces between 43 vegetated patches and the free-stream, spawning coherent vortices and eddies (Nepf 1999; White and Nepf 2008; Zong and Nepf 2010). Although it has been found in some cases (e.g., Wilkerson 44 45 2007; Hopkinson and Wynn 2009) that plant-flow interactions are similar on banks and 46 floodplains, and that turbulence levels on inclined non-vegetated and sparsely-vegetated 47 streambanks are sometimes similar (Hopkinson and Wynn 2009), other studies (Nepf 1999; 48 McBride et al. 2007) document elevated turbulence levels within sparsely-treed floodplains 49 relative to the non-vegetated case. This apparent dichotomy suggests that site-specific 50 conditions, such as the angle of the bank face and/or bank toe, may influence the relationship 51 between vegetation and channel hydraulics (McBride et al. 2007; Wilkerson 2007). Therefore, 52 the objective of the present study is to use a Froude-scaled flume experiment to characterize the 53 interacting influences of bank angle and vegetation density, defined as the ratio of plant frontal 54 area (the area of submerged leaves and stems in a vertical plane perpendicular to the channel 55 centerline) to flow area, on near-bank patterns of shear stress and turbulence.

56 Methods

Experiments were conducted in a 6.0 m \times 0.6 m \times 0.6 m recirculating flume set at a fixed 57 slope of 0.01 m m⁻¹. At the inlet, a rock-filled baffle box and 0.30 m-long, 0.02 m-diameter tubes 58 59 (flow straighteners) were used to dampen turbulence and provide parallel streamlines. To 60 simulate a sloping bank surface along one side of the flume, a 4.88 m-long insert, inclined at either 15° or 30° from the horizontal, was installed immediately downstream of the flow 61 straighteners. Stands of artificial vegetation of two different stem densities (defined as the 62 63 number of stems per square meter of bank surface) and two different leaf scenarios ("leaved" and "leafless"; the prefix L is used throughout the text to denote leaved cases; e.g., LD_{lo}, LD_{hi}) were 64 65 installed in a staggered pattern on the bank surface (Figure 1). Stems for the artificial plants were 66 constructed using acrylic rods; ten 28-gauge wire "branches" with 25×35 mm flexible "leaves"

made of contact paper were affixed to the rods in a pattern similar to Wilson et al. (2006a) and commencing 60 mm above the base of the stem. Other than the vegetative elements, the boundaries of the flume were smooth: the flume walls were constructed of lacquered marine plywood. Flow depth was controlled by a weir at the outlet, creating a gradually-varied, highlysubcritical (Table 1), and fully turbulent (15644 $\leq Re \leq$ 16095) flow field. Water depths were always less than the height of the plants and hence the plants were emergent.

73 Flume geometry was Froude-scaled from the Goodwin Creek bendway site in North 74 Mississippi (Langendoen and Simon 2008; Simon et al. 2000; Simon and Collison 2002; Wood 75 et al. 2001) to establish both geometric and kinematic similitude (Table 1). Bank slope lengths 76 and angles were computed for repeat surveys at eleven cross-sections at Goodwin Creek, and 77 two (cross-sections 5 and 6; Table 2) were selected for representation in the flume. The 15° and 30° bank angles in our physical model approximate the 10th and 90th percentiles observed in the 78 79 prototype, respectively. Owing to limitations imposed by the dimensions of the flume, the 80 selected length scales were computed using the ratios between the model slope lengths (0.41 and 0.46 m, respectively) and the 16^{th} percentile slope length (~2.0 m), rather than the median slope 81 82 length (3.3 m) of the prototype bank. This scaling approach yielded mean Froude scaling factors of 4.88 and 4.35 for the 15° and 30° bank surface, respectively. 83

84 Features (i.e., stem diameter, stem density, frontal area and flexural rigidity) of the 85 artificial vegetation were also scaled (Table 1). Vegetation models were based on willow and cottonwood yearlings up to 2 m-tall and 20 mm-diameter, which are commonly found on 86 periodically-inundated bank surfaces in densities of ~ 10 to 30 stems m⁻² (Wilson et al. 2006b). 87 Thus, applying Froude scaling, artificial plants were constructed using 450 mm-long, 4.54 mm-88 diameter acrylic rods and arranged with stem densities of 202 and 615 stems m⁻², respectively, in 89 90 a 3 m-long array, beginning immediately downstream of the flow straighteners. The flexural 91 rigidity (J) of stems was also Froude-scaled (Table 1), based on field data collected during the 92 present study (see Czarnomski 2010 for further details) and values reported by others (Niklas 93 1992; Freeman et al. 2000; Wilson et al. 2003). Note that Reynolds number similarity was 94 necessarily relaxed (Yalin 1971).

95 *Near-bed velocity measurements*

96 Near-bed velocities were measured at seven cross-sections spaced 0.055 m apart at 97 approximately 5 mm above the bed. In order to limit the influence of conditions imposed at the 98 inlet and outlet, cross-sections were located 1.84 - 2.23 m downstream from the flow straighteners. Velocities were measured at 25 Hz for 300 seconds with a downward-looking 10 99 100 MHz Nortek acoustic Doppler velocimeter (ADV) that was aligned with the z-axis (see Figure 1 101 for a definition of the Cartesian coordinate system employed). Sampling frequency was selected 102 assuming a Strouhal number of 0.21 (e.g., Schlichting 1968), estimating the likely eddy shedding 103 frequency caused by model stems (9.1 - 12.3 Hz) and then considering the Nyquist sampling 104 theorem. Sampling duration was selected after analysis of the cumulative velocity variance 105 associated with different sampling windows (e.g., Sukhodolov and Rhoads 2001). The sampling volume of the ADV had a diameter of 6 mm and a volume of 254 mm³, thus capturing turbulent 106 107 eddies that were approximately as small as the stem diameter. Boundary measurements were 108 made at fixed X-V coordinates for each of the seven cross-sections and were generally located 109 0.01- 0.03 m away from the nearest stem. However, if a velocity sampling location fell within 110 0.01 m of a stem, that stem was temporarily removed to permit data acquisition. ADV data with average correlation coefficients < 0.6 and signal-to-noise ratios < 0.15 dB were removed and the 111 112 remaining data were despiked using the phase-space threshold algorithm (Goring and Nikora 113 2002) within WinADV version 2.027 (Wahl 2009).

114 *Analysis of velocity measurements*

115 Using near-bed velocities measured at the cross-section 2.0 m from the beginning of the 116 vegetation, we computed the root mean square (RMS) difference between the instantaneous 117 velocities (U, V, and W) in the streamwise (X), lateral (V), and vertical (Z) directions and their 118 respective time-averages $(\bar{u}, \bar{v}, \text{ and } \bar{w})$ to represent turbulence intensity (Hinze 1975) and to 119 provide an indication of where shear stresses are highest (Biron et al. 2004; Hopkinson and 120 Wynne 2009). Computed values of *RMS* were normalized by the cross-sectional mean velocity 121 (U) to facilitate comparison of the three components and to illustrate the magnitude of turbulent 122 fluctuations relative to the mean flow.

Local estimates of lateral and vertical Reynolds stresses (τ_{uv} and τ_{uw} , respectively) were used as proxies for applied shear stress (e.g., Biron et al. 2004) and to quantify the magnitude and direction of turbulent fluctuations that represent momentum exchange across a given plane 126 (Robert 2003). τ_{uv} and τ_{uw} were estimated for all sampling points using $\tau_{uv} = -\rho \overline{u'v'}$ and $\tau_{uw} = -\rho \overline{u'w'}$, respectively, where primes denote fluctuations about the time averaged velocities.

128

129 Results

130 *Relative turbulence intensity (RMS/U)*

131 The presence of vegetation on the bank surface generally increased RMS/U at the base of 132 the bank slope and immediately adjacent to the bank (Figure 2). For example, at the base of the 133 bank slope, *RMS/U* increased by 120 - 650% over the non-vegetated scenario for LD_{lo}, LD_{hi} and 134 D_{hi}. At this location, values of *RMS/U* were also much higher for leaved than for leafless 135 vegetation. For example, relative to the non-vegetated case, at the base of the 15° slope, RMS_{μ}/U , 136 RMS_{V}/U , and RMS_{W}/U increased by 60 – 150% during leafless vegetated runs, but by 220 – 320% 137 during leaved runs (Figure 2). A similar result was true for the 30° bank: at the base of the bank 138 slope, *RMS/U* increased by 140 - 220% during leafless runs and by 350 - 650% during leaved 139 runs (Figure 2).

Streamwise relative turbulence intensity ($RMS_{u'}/U$) ranged from 85 – 160% of the lateral relative turbulence intensity ($RMS_{v'}/U$), and from 210 – 490% of the vertical relative turbulence intensity ($RMS_{w'}/U$). The differences in intensity were similar for the 15° and 30° bank slopes, although the peak magnitude of $RMS_{u'}/U$ on the 30° bank was up to 30% larger than $RMS_{u'}/U$ on the 15° bank and $RMS_{v'}/U$ was 100% larger on the 30° bank than $RMS_{v'}/U$ on the 15° bank.

145 *Reynolds stresses*

146 Spatial patterns of near-bed values of τ_{uw} and τ_{uv} were similar to patterns of *RMS/U*, 147 where values were generally positive and increases in stress were observed with increasing plant 148 density. Without vegetation, τ_{uw} values were generally positive (0 – 0.05 Pa) and were mostly 149 distributed uniformly throughout the cross-section (Figure 3). Once vegetation was introduced, 150 values of τ_{ijw} were positive within the main channel, with a local maximum near the center of the 151 main channel, and negative on the bank surface, with a local minimum near the base of the slope 152 (Figure 3). τ_{UW} was up to an order of magnitude lower for the 15° bank than the 30° bank, and the magnitude of τ_{UW} at the stationary points (e.g. maxima, minima) increased with increasing 153 154 vegetation density (Figure 3). Similar patterns were observed for τ_{uv} , where increasing vegetation 155 led to higher values of τ_{uv} , indicating increases in lateral momentum exchange across the base of 156 the slope. However, the magnitude of τ_{uv} across the slope base was similar along the 15° and 30° 157 banks (Figure 4).

158 The dominant orientation of stresses and momentum exchanges was more variable on the 159 15° bank than the 30° bank (Figures 3 and 4). For the 30° bank, lateral momentum exchange was 160 the primary stress found throughout the channel when vegetation was not present. With high 161 density vegetation along the 30° bank, the primary stress on the bank surface was τ_{uw} , while τ_{uv} 162 was higher along the base of the slope and in the main channel. For the 15° bank, τ_{uv} was 163 dominant at the base of the slope when no vegetation was present, but neither τ_{uv} nor τ_{uw} was 164 consistently dominant when vegetation was present.

165 Summary and Conclusions

166 This paper has presented results from an experimental study aimed at characterizing the 167 influence of bank angle and vegetation density on near-bank patterns of shear stress and 168 turbulence. The key findings of the study are:

- Increasing bank angle caused increased turbulence intensities and Reynolds stresses
 at the base of the bank slope. However, on the bank slope itself, relative turbulence
 intensities and Reynolds stresses were insensitive to the angle of the bank;
- Increasing vegetation density on the bank surface caused increased near-bed turbulence intensities and Reynolds stresses in the main channel and at the base of the slope. These increases were particularly evident along the base of the slope, supporting the findings of previous studies (e.g., Yang et al. 2007; Gorrick 2009; Hopkinson and Wynn 2009); and
- Relative turbulence intensities and Reynolds stresses were higher for leaved than for
 leafless conditions. This result highlights the importance of including leaves or
 equivalent canopy roughness in both flume and numerical experiments and casts
 doubt on the results of studies that have not done so. The additional frontal area
 afforded by a canopy, and the hydraulic behavior of a canopy, cannot be replicated by
 merely increasing stem density, but instead requires the use of vegetative elements of
 a more realistic morphology (e.g., Yang et al. 2007; Hopkinson and Wynn 2009).

184 It is acknowledged that the strength of these conclusions may be reduced by the lack of uniform 185 flow in our flume and we therefore encourage future studies to more carefully develop uniform 186 flow conditions (Tracy and Lester 1961). Nonetheless, the results presented herein contribute to the growing knowledge (e.g., Nepf 1999; Wilson et al. 2003; McBride et al. 2007; Yang et al. 2007; Gorrick 2009; Hopkinson and Wynn 2009) of the influence of vegetation morphology and configuration on near-boundary hydraulics. Furthermore, they emphasize the need to consider the morphology of vegetation when assessing turbulence and stress within patches of vegetation, and to evaluate the timing of flood events relative to leaf-out when planting vegetation as a management strategy to deflect near-bank flows.

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Table 1. Flume and prototype scaling parameters. Hydraulic parameters are presented as means for all scenarios. The scaling factor, λ , is 4.88 for the 15° bank surface and 4.35 for the 30° bank surface. Froude scaling relations are given by Julien (2002) and Wilson et al. (2003).

	scenario	scaling relation	model channel	prototype channel
Geometry				
bank slope length	15°	λ-1	0.41	2.0
(m)	30°	λ^{-1}	0.46	2.0
vertical bank face height	15°	λ^{-1}	0.49	2.77
(m)	30°	λ-1	0.37	2.77
Hydraulics	1.50	a -1	0.00	1.05
main channel flow depth	15°	λ 1	0.38	1.85
(m)	30°	λ^{-1}	0.42	1.85
cross-sectional mean velocity ¹	15°	$\lambda^{1/2}$	0.19	0.40
$(m s^{-1})$	30°	$\lambda^{1/2}$	0.21	0.47
cross-sectional mean Fr	15°	λ^{0}	0.13-0.16	0.13-0.16
(-)	30°	λ^{0}	0.15-0.20	0.15-0.20
Vegetation				
stem density	15°	λ-2	202	10
(stems m^{-2})	30°	λ-2	615	30
flexural rigidity	15°	λ-5	0.0435	120.4
$(N m^2)$	30°	λ^{-5}	0.0435	67.8
vegetation density	D _{lo}	λ^{0}	0.027; 0.029	0.027; 0.029
(15°; 30° bank surface)	D_{hi}	λ^{0}	0.085; 0.103	0.085; 0.103
	LD _{lo}	λ^{0}	0.155; 0.192	0.155; 0.192
	LD _{hi}	λ^{0}	0.468; 0.586	0.468; 0.586

¹ Velocity in the prototype channel was estimated based on the length scale factor and channel dimensions.

Table 2. Summary statistics of 64 evaluations of bank slope length and bank slope angle estimated from surveys of cross-sections 5 and 6 at the Goodwin Creek bendway, MS., established at this site in February 1996 and resurveyed at regular intervals until May 2003. The bend apex was initially at cross-section 4 and gradually migrated downstream to between cross-sections 7 and 8.

Statistic	Bank slope length (m)	Bank slope angle (°)
Minimum	1.00	8.8
10 th Percentile	1.71	15.7
16 th Percentile	1.98	17.3
Median	3.29	22.8
84 th Percentile	4.63	28.0
90 th Percentile	4.89	29.4
Maximum	5.61	39.8





y/b

y/b





FIGURE CAPTION LIST

Figure 1. Experimental design and flume cross-sectional design. D_{lo} is low density, no leaves; D_{hi} is high density, no leaves; LD_{lo} is low density, with leaves; and LD_{hi} is high density, with leaves.

Figure 2. Cross-stream variations of near-boundary relative turbulence intensity (*RMS/U*) in the *u* (longitudinal), ν (transverse), and w (vertical) directions for the 15° (a, b, c) and 30° (d, e, f) bank surfaces. Velocities were measured 2.0 m downstream from the beginning of the vegetation. Cross-stream position (*y*) has been normalized by channel width (*b*). D_{lo}, D_{hi}, LD_{lo} and LD_{hi} are defined in Figure 1.

Figure 3. Spatial patterns of vertical Reynolds stress (τ_{UW} , Pa) for the a) 15° and b) 30° bank surfaces. D_{lo}, D_{hi}, LD_{lo} and LD_{hi} are defined in Figure 1. The solid line represents the base of the bank slope.

Figure 4. Spatial patterns of lateral Reynolds stress (τ_{uv} , Pa) for the a) 15° and b) 30° bank surfaces. D_{lo}, D_{hi}, LD_{lo} and LD_{hi} are defined in Figure 1. The solid line represents the base of the bank slope.